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GROUND-WATER FLOW ALONG THE TOTAL SHORELINE OF AUSTIN LAKE AND ITS ENVIRONMENTAL CONTRIBUTION TO POLLUTION OF THE LAKE

by

William Thomas Williams

A Dissertation
Submitted to the Submitted to the Faculty of The Graduate College in partial fulfillment of the requirements for the Degree of Doctor of Philosophy Science Education

Western Michigan University Kalamazoo, Michigan April 1982

GROUND-WATER FLOW ALONG THE TOTAL SHORELINE OF AUSTIN LAKE AND ITS ENVIRONMENTAL CONTRIBUTION TO POLLUTION OF THE LAKE

William Thomas Williams, Ph.D. Western Michigan University, 1982

The vertical direction of ground-water flow in Austin Lake was determined by placing minipiezometers along the entire shoreline at fifty randomly selected sites. Groundwater samples were taken at each site and the specific conductivity of the water samples was determined along with the chloride and phosphate content. The data were analyzed to determine if there were any relationships among water movement, chemical content, and the presence or absence of an onshore septic tank. The minipiezometer data show that the primary direction of water movement through the sediment-water interface is downward but may become reversed at localized regions as conditions vary. The negative hydraulic differential is greater at the north and south ends of the lake than on the east and west sides. This observation makes a stronger case for reports by other researchers that Austin Lake is on the ground-water divide and that ground water flows to the north and to the south from the lake. Computer analyses of all data show relationships among specific conductivity, chloride content, and the presence of a septic tank. Offshore from septic tanks, the specific conductivity and chloride content

vary more and have higher average concentrations than offshore from areas without a septic tank, although visual inspection of the data shows pollution to be minimal. Phosphate content failed to show relationships with septic tanks, but tended to be higher on the west side of the lake— an area bounded by outwash and containing older functional and non-functional septic tanks.

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Lastly, I want to thank my wife, Eva, for her patience and support she has given me throughout it all.

William Thomas Williams

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CHAPTER 1

INTRODUCTION

Geographical Location

Austin Lake is located in the southeastern corner of Portage Township, Kalamazoo County, Michigan (Pigure 1). The lake encompasses parts of sections 23, 24, 25, 26, 35, and 36, T. 3 S., R. 11 W.

Geology

Austin Lake lies in a region glaciated by a succession of at least four major continental ice sheets. The main topographical features of the area were formed during the recession of the last ice sheet. Austin Lake is a "kettle" lake, occupying a depression formed from the melting of an isolated block, or blocks, of ice left during the retreats and readvances of the Lake Michigan lobe of the glacier during the last Ice Age— Wisconsinan time (Deutsch, Vanlier, & Giroux, 1960). The three-dimensional shape of the basins suggests that the lake was formed by the melting of three blocks of. ice that were partly connected (Straw, 1978). Organic sediments have filled the lake basins until at the present time the maximal surface water depth is less than 10 feet (3-0 m). The

Figure 1. Location of study site in Kalamazoo County, Michigan.

basins formed by the three blocks of ice range in depth from 25 feet (7.6 m) to 50 feet (15.2 m) with the deepest being in the southwestern portion of the lake (Figure 2). Below the organic material in the basins, the sediment varies from coarse gravel to clay (Figure 3). Along most of the shoreline, the sediments have been modified by waves so that sand and gravel predominate.

Surface glacial deposits surrounding Austin Lake are of three types--morainal, outwash, and channel deposits. Adjacent to the lake on the northeast side are morainal deposits that trend northeastward. These morainal deposits were deposited directly from glacial ice with running water playing a minimal role in deposition. The west and south sides consist of outwash--material that was deposited from meltwater streams issuing from glacial ice. On the east side of the lake, trending southeastward, and from the north part of the lake northwestward are channel deposits. These are similar in origin to the outwash deposits, but have been reworked and covered with finer materials (Deutsch et al., **¹⁹⁶⁰**).

Sources of Water Entering and Leaving Austin Lake

Austin Lake covers an area of 1050 acres (2593-5 ha), but drains a land area that encompasses only 340

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Figure 3. Basal sediment in Austin Lake (Straw, 1978).

acres (839.8 ha) (Straw, 1978). The effective drainage . basin is considerably smaller because the permeability of the surrounding soil is high (Figure 4). Some surface water enters the lake from storm sewers, but otherwise surface water from the immediate drainage basin contributes little to .the water budget. Some small ephemeral streams are present in areas of moderate to steep slopes, but most such drains do not contribute water unless rainfall is heavy. Other sources of water include canals from Long Lake and West Lake, direct precipitation, the Upjohn recharge pond, surface condensation, and ground-water discharge.

An outlet at the southeastern end of the lake leads to Gourdneck Creek. Water also seeps through the bottom to become part of the ground-water flow. Some water is lost through evaporation from the free water surface and through irrigation (Figure 5).

Austin Lake, with the other lakes in the complex, are all considered to be part of the St. Joseph River basin (Figure 6). Yet, Allen, Miller, and Wood (1972) placed the ground-water divide (major deep ground-water flow) on an east-west line through the center of Austin Lake with ground-water movement to the north and south of the lake (Figure 11).

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Figure 4. Drainage basin of Austin Lake (Straw, 1978).

Figure 5. Water sources and losses of Austin Lake.

Figure 6. The primary drainage basins of Kalamazoo County, Michigan (Allen et al., 1972).

Location of Sewer Lines

The dwellings around Austin Lake are served by sewer lines with the exception of those in the southwestern corner. Because sewer lines surrounding the lake are of recent origin many residents still use septic tanks.

History

Natural accumulations of organic sediments have partly filled the lake since the retreat of the last glacier. Radiocarbon dating of the organic sediments in certain parts of the lake indicate that the rate of accumulation of these materials was about 1 inch (2.54- cm) every 20 years for the period 8000 B. P. (before present) to 3000 B. P., 1 inch (2.54 cm) every 12 years for the period 3000 B. P. to 1700 B. P., and 1 inch (2.54- cm) every 7 years since 1700 B. P. It is likely that sedimentation may now be taking place at the rate of 1 inch (2.54- cm) every 2 years (Straw, 1978).

In 1962 an attempt was made to pump out the organic sediment near the north end of the lake. The objectives of the project were to determine (1) how much solid material could be pumped out in one hour using a four inch pipe, (2) if the hole formed by the pumping of muck would be filled in by the surrounding material, and (3) at what rate the hole would fill (Zeno, Kalamazoo

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Gazette, September 26, 1962). A holding area was constructed and filled with the pumped muck. The holding area was permeable so the water could drain out leaving the solids behind. After being filled, the weather turned cold, freezing and damaging the equipment. No further action was taken (Ryskamp, Note 1).

Also in 1962, concern over the lake level led to a plan for a ponding project to hold water to be released when the level of the lakes dropped (Kalamazoo Gazette, December 4, 1962). In 1963 attention was directed towards pumping well water into the lakes to maintain their levels (Kalamazoo Gazette, Swptember 21, 1963). In 1964 a proposal was presented to divert water from the Upjohn pond to Austin Lake (Kalamazoo Gazette, February 23, 1964). This became known as the Tri-Lakes Upjohn Project and, after a series of legal battles, construction finally began in December of 1966 (Kalamazoo Gazette, December 23, 1966). On July 20, 1967, (Kalamazoo Gazette) the pipeline construction was completed and five million gallons of water a day began flowing into Austin Lake.

The pipeline was instrumental in solving the low water-level problem, but because the lake is shallow, sunlight is able to penetrate to the bottom stimulating the growth of rooted aquatic plants. Changes in the use

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of the lake from predominantly fishing to fishing, boating, and water skiing caused residents to view the growth of weeds with concern. Various methods have been proposed to alleviate this problem. These methods include dredging, aeration, chemical treatment, and harvesting, but costs and uncertainity of the results have prevented any action. More basic research to reduce the degree of uncertainity needs to be done before major steps can be taken to improve the lake basin.

Assessment of Lake Environments

This paper is a result of an investigation into the movement and quality of ground water as it relates to Austin Lake. The purpose of the investigation was to locate areas of inflow and outflow near the shoreline of Austin Lake and to determine if the lake is being polluted from ground-water sources.

The main hydrogeologic factors that are prerequisites for assessing lake environments are: (1) regime dominance (the relative magnitude of ground water in the water budget of a lake); (2) system efficiency (a term describing the rate aspects of surface- and ground-water movement through a lake system); and (3) position within a ground-water flow system (Born, Smith, & Stephenson, 1974). This paper is concerned with this last factor

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along with the water quality of the ground water that enters the lake.

The best information for determining the position of a lake in a ground-water flow system is the groundwater potential distribution beneath the lake bottom, and the vertical hydraulic gradient is the ideal criterion. If the vertical gradient is positive (potential increasing downward) the lake is in a ground-water discharge region. If the vertical gradient is negative (potential decreasing downward) the lake is in a groundwater recharge region. If the vertical gradient is zero or negative at one end and positive at the other, the lake is in a region of lateral flow. Measurement of this factor must be done onsite with piezometers (Born et al., 1974).

Significance of Study

This study is significant for three reasons: (1) little research has been done on the ground-water flow beneath lakes in Michigan and none on Austin Lake with the exception of generalized deep ground-water flow patterns. The ground-water flow into and out of the lake has not been determined in detail by use of piezometers; (2) the lakeshore is highly urbanized and there is the possibility of substantial pollution; and (3) a detailed

knowledge of the hydrology of Austin Lake must be gained before an adequate plan can be developed for the restoration of the lake.

CHAPTER 2

REVIEW OP PERTINENT LITERATURE

Ground-water flow and its association with lakes has only recently become the object of detailed research studies. Because of this, information on the subject is limited. Research on the quality of ground water that enters lakes is virtually non-existent. This chapter reviews research and writings of ground-water flow as it relates to lakes.

Ground-water Plow

Preeze and Cherry (1979) described ground-water movement in homogenous and isotropic materials as a flow from highlands towards valleys. In this situation the water table forms a subdued replica of the topography. The symmetry of the system creates vertical boundaries beneath the valleys and ridges across which there is no flow. These "ground-water divides" usually coincide exactly with surface-water divides, and their orientation is, for all practical purposes, vertical (Pigure 7). Preeze and Weatherspoon (1966) showed that in more complex topographic and hydrogeologic environments groundwater movement becomes more complicated.

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Toth (1963) and Born et al. (1974) described groundwater flow as ground-water movement in response to gravity from regions of high to low potential along flow lines that describe the path of an individual waiter particle as it moves through the soil. A set of flow lines oriented so that any two flow lines adjacent at one point © in the system remain adjacent throughout the system make up a ground-water flow system. All the flow lines in a ground-water flow system can be intersected by an uninterrupted surface across which flow takes place in one direction only. Freeze and Cherry (1979) stated that flow lines deliver ground water from recharge regions to discharge regions (Figure 8).

Breeze and Cherry (1979) described ground-water flow beneath a hilly topography. In this situation there are numerous subsystems within a major flow system. Water that enters the flow system in a given recharge region may be discharged in the nearest topographic low. (local ground-water flow) or it may be transmitted to the regional discharge region in the bottom of the major valley. Toth (1963) showed that as the depth to lateral extent of the entire system becomes smaller and as the amplitude of the hummocks become larger, the local systems are more likely to reach the basal boundary, creating a series of small independent cells (Figure 9)•

Effect of topography on regional ground-water
flow patterns (Freeze and Cherry, 1979). Figure 9.

 \cdot
Preeze and Cherry (1979) listed the effects geological heterogeneity can have on regional ground-water flow. Heterogeneity can affect the interrelationship among local and regional ground-water flow, the surficial pattern of recharge and discharge regions, and the quantities of flow that are discharged through the systems.

Born et al (1974) discussed lakes and their associa-tion with ground water. Lakes situated in ground-water recharge regions (recharge lakes) contribute to the ground water throughout the entire lake bottom. Lakes in ground-water discharge regions (discharge lakes) gain ground water throughout the entire lake bottom. In regions of lateral ground-water flow, lakes lose to the ground water on one side and gain ground water on the other side. Such lakes are known as flow-through lakes (Pigure 10). Lakes may be associated with the regional ground-water flow, local ground-water flow, or both. It is possible for lakes to intersect both a shallow and a deep flow system with the result that ground-water movement near the lake margins may be opposite to the directions of movement near the lake center.

Manson, Schwartz, and Allred (1968) indicated that "muck" is highly impermeable and therefore, where it exists, prevents water from moving in and out of a lake. They stressed the importance of bottom materials in preventing or permitting ground water to enter or to leave

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Figure 10. Configuration of ground-water
flow systems around lakes
(Born et al., 1974)

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lakes. They also speculated that glacial lakes--not intersecting the water table— upon initial formation did not hold water. Finer materials carried in by sheet and gully erosion sealed the bottom of the kettles, thereby permitting the lakes to hold water. They presented data to support the view that the permeability of the soil and drift adjacent to the pond or lake are usually much greater than that of the bottom. Thus, when the water level rises to flood the normal basin, infiltration is rapid and the level tends to decline rather promptly to the usual level. Their investigation of ponds and small lakes in relatively deep depressions showed that inflow underground was largely limited to wet periods— and even then was not prevalent at all sites.

McBride and Pfannkuch (1975) and Lee (1972) showed that seepage into or out of lakes tends to be concentrated near the shore and decreases with increasing distance from the shore. In many places the rate of decrease is exponential. Their model indicated that little of the ground-water flow passes beneath the lake. Instead, most of it is diverted upward as it approaches the lake, passes through the lake, moves downward, and eventually comes back to the horizontal. McBride and Pfannkuch (1975) showed that this holds true only where the diameter of the lake is at least roughly comparable to the thickness of the underlying permeable materials.

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Where ponds have diameters much less than the thickness of the sections, seepage occurs almost uniformly across the bottoms of the depressions.

McBride and Pfannkuch (1975) in citing Lee's (1972) study of Lake Sallie, Minnesota, noted that in the 2 km wide lake, one-half of the discharge occurs within 17 m of shore, **90** percent within 60 m and 99 percent within 120 m. The edge of the fine-grained sediments in the center of Lake Sallie is 300 m from shore. This and similar research done by Hackbarth (1968) in the Little St. Germain Lake Basin in Wisconsin show the importance of the geometry of the ground-water system and the relatively small role that lake sediments have in controlling the distribution of seepage in lakes.

Deutsch et al. (1960) and Allen et al. (1972), using water level data from observation wells, lakes, and streams, determined the direction of the regional groundwater flow in the Austin Lake complex. This complex includes Austin, West, Long, Sugarloaf, and Gourdneck lakes. These data indicated that the ground-water movement was both north and south of an east-west line crossing the approximate center of Austin Lake (Figure 11).

Allen et al. (1972) devised a water budget for the complex from October 1, 1965 to March 30, 1966 and again from April 1, 1966 to September 30, 1966. For their first budget, they cited a subsurface inflow of 73 acre-feet

Figure /1/1. Ground-water divides as shown by Deutsch et al. (1960) and Allen et al. (1972).

and a subsurface outflow of 1.070 acre-feet. For the second budget, the subsurface inflow is 146 acre-feet and the subsurface outflow is 1,010 acre-feet. They said the subsurface outflow returns to the surface at Gourdneck Creek, a short distance to the south of Austin Lake and flowing east. These figures indicate that little ground water enters Austin Lake as compared with the total flow.

Ground-water Quality

Legrand (1965) described contaminant movement from waste sites into the subsurface water circulation. This movement results in contaminated zones, or enclaves, in the zone of saturation. The contaminants tend to be entrained in ground-water flow and also tend to attenuate to varying degrees by dilution in water, decay with time, or some other 'die away' mechanism, and adsorption on earth materials. Once a contaminated zone is approximately stable, an increase in concentration may cause it to remain about the same size under some conditions and enlarge in others according to combined attenuation effects. Where attenuation occurs only by dilution the contaminated zone will become enlarged with increased concentration, and even if attenuated only slightly by dilution may become greatly elongated in the direction of ground-water flow. Where attenuation occurs through

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decay or adsorption or through both mechanisms, the contaminated zone may not change appreciably, even if the concentration of contaminants increases. Where the contaminants sink into the ground, the vertical direction predominates in the zone of aeration that lies above the water table. When they reach the zone of saturation contaminants tend to spread laterally (Figure 12).

Lee (1972) , in his study of lakes in a glacial terrain, found that septic tank effluent from lakeside houses entered the lake within 9 meters of the shoreline in a fan-shaped pattern. His study showed that 40 percent of the nitrogen entered the lake although the phosphorus was fixed in the soil.

Previous studies of the water quality of Austin Lake indicate low concentrations of dissolved nutrients. Jones and Henry Engineers, Limited (1972) showed that chloride concentrations on the western edge of the lake averaged 56 mg/1 and 20 mg/1 at other sampling points. Orthophosphates ranged from 0 to 0.45 mg/1.

Allen et al. (1972) listed chloride concentrations between 1 and 6 mg/1.

Chemical analyses by the author during a period from September to December, 1977, show low levels of phosphorus.

Information on the quality of ground water discharged into Austin Lake is not available.

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CHAPTER 5

PROCEDURE FOR COLLECTING DATA

The purpose of this investigation was twofold. The first objective was to determine the direction of water flow through the sediment-water interface along the total shoreline of Austin Lake. The second objective was to determine if this water was contributing to the degradation of the lake.

Determination of Hydraulic Head

The total shoreline of Austin Lake was divided into segments. Most lots represented one segment. Large lots were divided into two or more segments. Areas not designated as lots were divided into segments that were approximately 50 feet in width. Each of the 678 segments was assigned a number. Fifty segments were selected at random for the study (Figure 13)-

In order to determine the direction of ground-water flow, minipiezometers were placed at the approximate center point of each segment at distances of 10 feet (3.0 m) and 20 feet (6.1 m) from shore.

The minipiezometer, modified from one used by Lee and Cherry (1978), consists of a 3/4 inch (1.9 cm) iron

Figure 13. Approximate locations of sample sites
on Austin Lake.

pipe. A 1/2 inch (1.3 cm) iron pipe 8 inches (20.3 cm) long was welded to one end. A 7/16 inch (1.1 cm) polyethylene tube 8 inches (20.3 cm) long was placed inside the 1/2 inch (1.3 cm) iron pipe and the two were drilled simitaneously and perforated with a 1/4 inch (0.6 cm) bit. The polyethylene tube was then removed, wrapped with a fine mesh nylon net and replaced. Care was taken to be sure the holes on the pipe and the tube lined up. They were then fixed in place with a small bolt through the pipe and tube. The 1/2 inch (1.3 cm) pipe was capped with a reducer and plugged. The reducer and plug served as a driving point.

The pipe was capped and driven to the desired depth. The cap was then removed and water allowed to rise up into the pipe. To facilitate the process, water was added until the water level in the pipe was at the level of the water in the lake. Subsequent water-level movement within the pipe was measured and recorded along with water depth and pipe depth into the sediment. Waterlevel movement was checked by varying the.level of water in the pipe to determine if the results could be replicated. The difference between the water levels of the lake and pipe is the hydraulic head. To facilitate the measurement of the hydraulic head, a device (manometer) was inserted into the minipiezometer. Readings could then be made above the surface of the lake (Figure 14).

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Figure 14. Sampling equipment.

Water Sampling

The minipiezometer was pumped dry (not always possible in permeable sediments) and ground-water samples taken from water that refilled the pipe from below.

Chemical Analysis

Each sample was frozen at the end of each sampling day. When enough samples accumulated, they were thawed. The specific conductivity was determined and the sample analyzed for its chloride and phosphate content.

Conductivity was measured using model EP-/10 specific conductivity meter manufactured by the Myron L. Company of Encinitas, California. Results are given in micromhos/centimeter .

Chloride concentration was determined using the Hach mercuric nitrate method. In this procedure, diphenylcarbazone is the indicator and mercuric nitrate the titrant. Once all the chloride present in the sample is complexed by the mercuric ions, the excess mercuric ions combine with diphenylcarbazone to form a purple complex indicating the end point of the titration.

Phosphate concentration was determined using the Hach ascorbic acid method. In this phosphate analysis, orthophosphate reacts with molybdate in an acidic medium to produce phosphomolybdic acid that is then reduced to

a heteropoly blue compound. The phosphate concentration of the water sample is then determined by measuring the intensity of the blue color.

Importance of Water Quality Parameters Measured

Conductivity is a measure of the ability of water to conduct an electrical current. It is expressed in micromhos per centimeter at 25°C. Pure water has a low conductivity. Conductivity increases with increasing concentrations of dissolved minerals in the water. Normally the amount of dissolved solids in milligrams per liter (mg/1) is approximately 65 percent of the conductivity. Dissolved solids are often contributed to ground water by industrial and domestic wastes.

Soluble Orthophosphate— Phosphorus is the dissolved form of phosphate that is directly available for plant growth. Phosphate is reported rather than phosphorus be- . cause the element, phosphorus, does not occur free in nature. Instead, it is found in combined forms, the most common being phosphates. Phosphorus, like nitrogen, is an element essential to plant growth and both elements can be derived from sources such as soils and decomposition of organic matter. Phosphorus has become increasingly evident in the environment because of its past extensive use in detergents and its present use in agriculture. In streams, phosphorus is an extremely limiting

element for plant growth and thus an increased supply from outside sources such as sewage effluents may provide the stimulus for excessive production.

Phosphate movement in ground water is limited and therefore was not expected to be detected in large concentrations. Comparisons of phosphate concentrations with chloride concentrations may yield additional information regarding these species.

The chloride ion, in combination with one of several cations such as sodium, magnesium, and potassium, is a common impurity in most natural waters. Chloride compounds are readily dissolved from rocks by surface and subsurface waters. A concentration of several hundred milligrams per liter is necessary before a salty taste is imparted to the water or aquatic organisms are harmed.

Chlorides are essential in the diet and passes through the digestive system to become one of the major components of sewage. The wide use of salt in water softeners also contributes a large amount of chloride to sewage. In water, chloride concentrations are attenuated primarily by dilution and therefore high concentrations are good indicators of pollution.

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CHAPTER 4

ANALYSIS OP DATA

The data were analyzed on an IBM 3033 computer operating from an IBM 32?0 terminal under Time Sharing Option (TSO). The language used throughout the analysis was Statistical Analysis System (SAS).

Table Descriptions

The data analyzed are listed in Tables 1, 2, and 3.

Table 1, on pages 36, 37, 38, and 39, contains physical data from each sampling point. The data in Table 1 include the depth of the water at the sampling point (DW), the depth the piezometer was driven into the sediment (DP), the hydraulic head (HH), and whether or not a functioning septic tank was present on the lot onshore from the sampling site.

Table 2, on pages 40, 41, 42, and 43, contains the results of the ground-water sample analyses for each sampling point. The data in Table 2 include specific conductivity (COND), chloride content (CHL), and phosphate content (P).

Table 3, on pages 44, 45, and 46, contains data on other physical factors that may have an effect on groundwater movement and/or chemical properties of the ground

Physical Data for Each Study Site

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Table 1 (Continued)

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Table 1 (Continued)

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Table 2

Water Analysis Results for Each Study Site

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Table 2 (Continued)

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Table 2 (Continued)

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Location Number	Drainage Basin	$Soi1^a$	Glacial, Material ^b	Permeability ^C	
	Altitude (ft)			3 m	6 m
7	859	O ₂	0		3
18	861	O ₂	O.	2	2
32	864	02	\circ	1	1
49	868	03	O	2	S.
68	868	03	0	2	
75	868	03	0	1	1
120	876	01	O	3	1
144	868	01	0	2	3
168	872	03	O	3	3
170	870	03	O	2	2
176	860	03	O	2	3
185	861	03	O	2	2
192	870	03	0	2	2
213	868	04	O	3	3
218	868	03	O	3	3
244	869	04	O	3	3
260	868	04	O	3	3
274	864	04	0	3	3
275	864	04	0	2	3
281	863	04	O	3	3

Table 3 Factors Considered in the Statistical Analysis of Austin Lake's Ground-water Characteristics

Location Number	Drainage Basin Altitude (f _t)	${\tt soil}^{\tt a}$	$\mbox{Glacial}^b$ Material	Permeability ^C	
				3 m	6 m
287	862	04	\circ	3	3
305	858	A	C	2	2
320	861	B	C	2	\overline{c}
361	861	04	C	3	3
365	861	04	C	3	3
366	861	04	C	3	3
373	861	04	C	3	3
383	861	04	C	\overline{c}	2
401	861	04	C	3	3
403	861	04	C	3	\overline{c}
416	861	O4	C	3	3
419	861	04	C	3	3
420	861	04	C	3	3
451	859	04	C	3	3
457	859	04	C	3	3
473	859	04	M	3	3
496	861	04	M	2	2
497	861	04	M	2	2
508	861	04	M	2	2
519	861	04	M	3	2
548	858	04	\circ	2	2

Table 3 (Continued)

Location Number	Drainage Basin Altitude (f _t)	$Soi1^a$	Glacial _b Material	${\tt Permeability}^{\tt C}$	
				3 m	6 m
560	861	04	O	1	1
563	862	04	o		2
579	865	04	\circ	1	1
585	870	04	O	3	3
596	868	04	Ω	3	3°
614	859	Α	Ω	3	3
650	868	02	O	3	3
663	861	A	O	3	2

Table 3 (Continued)

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 α_{01} , 02, 03, 04 = varieties of Ostemo

A = Adrian

B = Brady

 $b_0 =$ outwash

M = morainal material C = channel deposits

 c_1 = excellent 2 = good 3 = poor

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water. These factors are the altitude of the drainage basin (DBA), the type of soil adjacent to and surrounding the lake, the surface glacial material adjacent to and surrounding the lake, and a subjective judgment of the permeability of the sediment at each sampling point.

The altitude of the drainage basin is the maximal altitude in the drainage basin perpendicular to the shoreline at the sampling site. The types of soil adjacent to and surrounding the lake include four varieties of Oshtemo (01, 02, 03, 04), Brady (B), and Adrian (A). The types of surface material adjacent to and surrounding the lake are channel deposits (C), outwash (0), and morainal material (M). Permeability is classified as excellent (1), good (2) , and poor (3) .

Variable Descriptions

Variables are classified according to whether they are measurement or classification variables and whether they are dependent or independent variables.

Measurement variables have measured values. Classification variables do not have measured values, or if they do, they can be simply grouped and given a group name. Hydraulic head, depth of the water, depth of the piezometer in the sediment, altitude of the drainage basin, specific conductivity, chloride content, and phosphate content are all measurement variables. The classification

variables are glacial material, soil, permeability of sediment, distance from shore, and presence or absence of a septic tank.

Dependent variables are subject to change if conditions change. Independent variables are fixed. They do not change. Hydraulic head, specific conductivity, chloride content, and phosphate content are dependent variables. Independent variables include all the classification variables.

Analytical Procedure

The first step in the analyses was to examine interdependencies among the variables DP, COND, HH, CHL, and P using pairwise Pearson Product Moment correlations (PROC CORR). In this procedure correlations were calculated considering breakdowns involving the independent variables— septic tank and distance from shore in the following manner: (1) all observations combined (Appendix A); (2) observations classified according to values on both independent variables (Appendix B); and (3) observations classified according to the presence or absence of a septic tank only (Appendix C).

The second step in the analyses used two-way analysis of variance procedures to determine the relationship between septic tank and distance from shore on the dependent variables HH, COND, CHL, and P. Because of unequal

sample sizes a general linear models program (GLM) was used for this analysis. This analysis was performed in the following manner: (1) observations classified according to values on both independent variables (Appendix D) and (2) observations classified according to the presence or absence of a septic tank only (Appendix E).

Schematic plots were drawn for each dependent variable to determine if visual inspection could identify any differences in the effects of a septic tank and/or distance from shore (Appendix E). Examination of the plots indicated that other factors might be related to the results. Factors considered were the altitude of the drainage basin, depth of water at each sampling point, the soil surrounding the lake, and the surface glacial material surrounding the lake. Those parameters were added to the data base along with a subjective determination of the permeability of the sediment at each sampling site (Table *3)•*

The dependent variable data were examined for skewness using the univariate procedure. Log transformations were made on the data to reduce the skewness for all dependent variables except hydraulic head (Appendix G).

In the final analyses the GLM procedure was used because both classification and measurement variables were among the independent variables. These analyses included the added factors and the log transformations.

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In the final analysis of HH, the classification variables were types of glacial material, types of soil, and permeability of the sediment. The measurement variables were the altitude of the drainage basin, depth of the piezometer, and the depth of the water (Appendix H).

In the final analysis of the other dependent variables (COND, CHL, P) the classification variables were septic tank, distance from shore, glacial material, soil, and permeability. The measurement variables were hydraulic head, altitude of the drainage basin, depth of the piezometer, and depth of the water (Appendix I).

In all analyses using the GLM procedure, the results referred to are from using the Type IV sum of squares. The Type IV sum of squares is the most appropriate for this study because it is adjusted for differences in sample size.

Computer Results

Hydraulic Head

In the initial correlational procedure observations were not grouped according to the independent variables septic tank and distance from shore. Here hydraulic head was seen to have a significant relationship to conductivity, r (84) = 0.21843, $p \le 0.0433$. At locations without a septic tank, there was a highly significant relationship

between hydraulic head and conductivity at a distance of 3 meters, r (25) = 0.49835, $p \le 0.0082$. At 6 meters, the relationship of hydraulic head to conductivity was marginally significant, r (24) = 0.36871, $p \le 0.0638$. At locations with a septic tank the relationship of hydraulic head to conductivity was marginally significant at 3 meters, r (14) = 0.43058, p \leq 0.0959. At 6 meters from shore, the relationship of hydraulic head to conductivity was not found to be significant, but the relationship of hydraulic head to phosphates was significant, r (**15**) = -0.48480 , $p \le 0.0486$. Because the relationship of hydraulic head to phosphates was shown to be significant in only one correlation set, it is likely that this statistic is the result of chance alone.

In the second series of correlations in which distance from shore was not a factor, the relationship of hydraulic head to conductivity was significant to localities where a septic tank was not present, r (**51**) = 0.34441, $p \le 0.0116$. When a septic tank was present, the relationship of hydraulic head to phosphates was marginally significant, r $(31) = -0.32440$, p ≤ 0.0655 .

The GLM procedure showed an overall significant relationship to hydraulic head, $F(3, 86) = 3.54$, $p \le 0.0179$. Individually there was a significant relationship of hydraulic head to the presence of a septic tank, $E(1, 86)$

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 $= 5.14$, $p \le 0.0259$, and a significant relationship of hydraulic head to distance from shore, $F(1, 86) = 4.31$. $p \le 0.0408$. The interaction between septic tank and distance from shore for hydraulic head was not found to be significant, $F(1, 86) = 0.41$, $p \le 0.5228$.

The schematic plot showed that the hydraulic head decreased with distance from shcre and decreased in the presence of a septic tank.

In the final analysis of hydraulic head, the independent variables were the class variables (glacial material, soil, and permeability) and the measurement variables (altitude of the drainage basin, water depth, and the depth of the piezometer). This model was highly significant, $F (12, 74) = 4.0, p \le 0.0001,$ and explained about 39 percent of the variability.

Overall, the relationship of hydraulic head to the type of glacial material was significant, $F(2, 74) = 3.87$. $p \le 0.0251$. The least squares means for channel deposits and outwash were -**52.8** and 49-5 respectively and the least squares mean for morainal material was -130.6. This indicates a greater outflow in the area of the lake that is bounded by morainal material.

The relationship of hydraulic head to soil type was marginally significant, $F (5, 74) = 2.10$, $p \le 0.0738$. Because of this overall marginal significance little emphasis is placed on it.

The relationship of hydraulic head to the permeability of the sediment was highly significant. $F(2, 74) = 9.62$. $p \leq 0.0002$. From the least squares means, poor permeability, with a least squares mean of -30.19, was associated with the highest hydraulic head (most positive number and lowest outward flow) as compared to excellent permeability with a least squares mean of -**159**.**0**.

The relationship of hydraulic head to the altitude of the drainage basin was significant, $F(1, 74) = 5.33$, $p \le 0.0237$. The estimate of the linear coefficient on the effect of the drainage basin altitude is -22.66. This is the slope of the linear regression line and being negative gives a result of a more negative head with an increase in the altitude of the basin.

The relationship of hydraulic head to the depth of the piezometer in the sediment was significant, $F(1, 74)$ $= 4.32$, $p \le 0.0411$. The estimate of the linear coefficient on the depth is -0.80. This means that the deeper the piezometer is placed in the sediment, the greater the negative hydraulic potential differential.

The relationship of hydraulic head to water depth was significant, $F(1, 74) = 4.29$, $p \le 0.0419$. The estimate of the linear coefficient on the water depth is -0.77- This is negative, therefore the greater the depth, the greater the negative hydraulic potential differential.

Conductivity

In the initial correlational procedure that did not group observations according to the independent variables septic tank and distance from shore, the correlation of conductivity with hydraulic head was significant as reported in the previous section. In addition, when no septic tank was present, conductivity had a significant relationship to chlorides at $\overline{3}$ meters, \underline{r} (27) = 0.38365, $p \le 0.0399$, and a highly significant relationship to chlorides at 6 meters, r $(26) = 0.53692$, p ≤ 0.0032 .

In the second series of correlations when distance from shore was not considered in the analysis, again conductivity had a significant relationship to hydraulic head as reported previously. Conductivity had a very high significant relationship to chlorides, r (55) = 0.43355, $p \le 0.0008$, when no septic tank was present. With the presence of a septic tank, conductivity had a marginally significant relationship to the depth of the piezometer, r (31) = -0.30350, $p \le 0.0860$.

The schematic plots for conductivity showed a marginal increase in the presence of a septic tank. There was a marginal increase with distance when the septic tank was absent and a decrease with the presence of a septic tank.

In the final analysis, the independent class variables

are glacial material, soil type, permeability, distance from shore, and septic tank. The independent measurement variables are hydraulic head, altitude of the drainage basin, depth of the piezometer in the sediment, and depth of the water.

The full model for the log of conductivity was not found to be significant, $F(15, 67) = 1.17$, $p \in 0.3151$. In this model conductivity had a marginally significant relationship to hydraulic head, F $(1, 67) = 3.19$, p \leq 0.0788. The estimate of the linear coefficient for the effect of the hydraulic head was 0.0004. This is positive and means that the greater the hydraulic head the greater the conductivity. In other works, the faster the water is flowing out of the lake, the lower the conductivity.

None of the other variables was found to be significant and these variables are not considered further.

A reduced model was designed for the log of conductivity. This model contained only the factors that showed significant values in the full model. These factors were hydraulic head and septic tank. This model was significant, F $(3, 82) = 3.29$, $p \le 0.0246$, and explained almost 11 percent of the variability. In this model conductivity had a significant relationship to hydraulic head, $F (1, 82) = 6.01$, $p \le 0.0164$. The presence of a septic tank no longer showed a particularly strong relationship, but when considered separately from hydraulic head, the
relationship became marginally significant. This implies that the conductivity is greater in the presence of a septic tank.

Chlorides

In the correlational procedures, significant relationships were not found other than that of conductivity previously mentioned.

The GLM procedure showed a marginal significant relationship between chloride content and septic tank presence, \underline{F} (1, 87) = 3.22, \underline{p} \leq 0.0761.

The schematic plot for chlorides showed a fairly strong relationship with the presence of septic tanks. Chloride content was greater and varied more in the presence of a septic tank.

In the final analysis, the independent variables were the same as in the analysis of the log of conductivity. The full model for the log of chlorides was not found to be significant, F (15, 68) = 1.53, p \leq 0.1198. Chlorides had a marginally significant relationship to permeability, $F(2, 68) = 2.38$, $p = 0.1001$, and a significant relationship to the depth of the water, \underline{F} (1, 68) = 4.13, \underline{p} \leq 0.04-59. The estimate of the linear coefficient for water depth was negative (-0.0016), therefore the deeper the water, the lower the level of chlorides.

None of the other relationships was strong enough to consider.

Phosphates

The initial correlations did not reveal any significant relationships between phosphate content and the several variables.

In the full model for the log of phosphates, the independent variables were the same as in the analyses of the other full models. The overall model was not found to be significant, F $(15, 66) = 1.27$, p ≤ 0.2450 . Glacial material had the only significant relationship to phosphate content, $E (2, 66) = 3.75$, $p \le 0.0288$. The least squares mean for outwash is 0.242 as compared to channel deposits at 0.074- and morainal deposits at 0.004-. This means that outwash has a greater level of phosphates, but overall, phosphate content was not found to be significant.

CHAPTER 5

DISCUSSION AND CONCLUSIONS

Ground-water Plow

The first objective of this investigation was to determine the direction of ground-water flow through the sediment-water interface along the total shoreline of Austin Lake. This was done by placing minipiezometers in the sediment beneath the water surface and measuring the hydraulic head. The data obtained show that at most sites the water movement has a downward component. The greatest downward flow appears to be at the northern, southern, and southwestern portions of the lake (Figure 15). Few measurements show ground water moving into the lake. The compass direction of flow cannot be determined from the data, but can be inferred from previous data as presented by Deutsch et al. (1960) and Allen et al. (1972). Their models suggest that ground water moves to the north and to the south from Austin Lake. Austin Lake is situated on the ground-water divide.

During the collection of hydraulic head data, three conditions were found to exist: (1) a negative hydraulic head for both samples at the sampling site; (2) a positive hydraulic head for both samples at the sampling site; and

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(3) a zero or a positive hydraulic head at 3 meters and a negative hydraulic head at 6 meters.

In the first situation, water is seeping through the lake seal and becoming part of the ground-water flow (recharge point). In the second situation, the ground-water movement is upward and into the lake (discharge point). The third situation suggests a combination of the first two. Ground water is moving into the lake at 3 meters and moving out of the lake at 6 meters (Figure 16). This last situation could occur if there is a slight local elevation of the water table near shore and would be enhanced by the presence of a channel for ground-water movement into the lake. These channels may be present. During sampling it was noticed that the sediment was layered at many points around the lake. These layers were narrow--from one to several centimeters in thickness. They were made of sand interbedded with finer material with organic material scattered throughout. At some of these sites the piezometer was driven deeper. The hydraulic head then became negative.

Computer analysis shows several factors affecting the hydraulic head. One factor is the depth of the piezometer in the sediment. This was expected. The deeper the piezometer is in the sediment, the more equipotential lines intersected, the greater the hydraulic head (Figure 8).

Another factor affecting the hydraulic head is the

Figure "16. Possible ground-water flow patterns at Austin Lake.

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depth of the water. As the depth of the water increases, the hydraulic head becomes more negative. This statistic is probably the result of the greatest outward flow being at the areas of the lake that were the deepest, namely the northern, southern, and southwestern portions.

The permeability of the sediment affects the hydraulic head. The largest hydraulic head differential occurs where the permeability is the greatest. At one sampling site that exhibited excellent permeability, a siphoning hose was placed from the lake into the minipiezometer. Water ran into the piezometer (out of the lake) for some time until the hose was removed. This indicates a thin seal in the lake. Once through the seal, the sediments are permeable.

The statistical analysis of the relationship of drainage basin altitude to hydraulic head is significant but not in the way expected. The higher the altitude of the drainage basin, the more negative the hydraulic head. Examination of the sampling results show high negative hydraulic head (outward) readings in the southwestern region of the lake. This region contains some of the highest points in the drainage basin. The southwestern region is illustrated on the cross-section of Austin Lake (Eigure 2). The left side of A-A' is the surface of Gourdneck Lake. This surface of Gourdneck Lake is below the surface of Austin Lake. It is reasonable to assume

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that this lower surface has a major effect on ground-water flow from Austin Lake towards Gourdneck Lake and that the .topographic high between the two lakes is overshadowed by this effect.

The northern shore of Austin Lake also shows a large negative head. The surface-water divide is to the north (Figure 6). This neans that surface water flows to the south and ground water flows to the north in this region. Therefore altitude of the drainage basin has a minimal effect on the total ground-water flow at Austin Lake.

In glacial sediment, morainal material shows the greatest outward flow (Figure 17). It is unclear whether this is a true picture of what is happening because of the small extent of morainal material adjoining the lake. These glacial materials are surface deposits and may have only a minimal effect on regional ground-water flow.

Pollution

The second objective was to determine if the lake is being polluted by ground water by examining the specific conductivity, chloride content, and phosphate content of ground-water samples from beneath the lake.

Three factors, hydraulic head, depth of the piezometer in the sediment, and presence or absence of a septic tank, affect the specific conductivity. The more positive the hydraulic head, the greater the conductivity (Figure 18).

Figure 17 Average hydraulic head readings for the different types of glacial material.

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In other words, the faster water is moving out of the lake, the lower the conductivity. The deeper the piezometer was placed in the sediment, the greater the conductivity. Conductivity is higher when a septic tank is present (Figure 19). Consequently, it appears that septic tank effluent is entering the lake. These results seem to indicate that at certain times during the year, at some points and under certain conditions, ground water moves into the lake bringing in dissolved compounds. Later when the ground-water direction is reversed, the compounds are carried downward and away from the lake, therefore the higher concentrations will be at the lower level. This idea is strengthened by the fact that a late summer check of the hydraulic head on the west side of Austin Lake indicated an outward flow even though some of the points had indicated an inward flow earlier in the spring and summer.

Chlorides exhibit some of the same relationships as conductivity. Offshore from septic tanks chloride content varies considerably. At locations without a septic tank chloride content varies little (Figure 20). This difference in variance indicates some inflow of effluent from septic tanks.

Significant relationships were not found between phosphates and any of the dependent variables. This is

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Figure **20. A** comparison of the chloride concentrations with the presence or absence of a septic tank.

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as expected as phosphates are rapidly attenuated in sediment. One interesting result is that a greater concentration of phosphates was present near glacial outwash. This area contains mostly older homes and the possibility exists that phosphates are entering the lake from older functional or non-functional septic systems (Figure 21).

Summary of Conclusions

The conclusions directly related to the objectives are:

- 1. The primary direction of water flow is downward through the sediment-water interface.
- 2. Under certain conditions ground water moves up through the sediment-water interface.
- 3. Upward movement of ground water brings in dissolved materials from onshore septic systems.
- 4. The amount of septic tank effluent entering the lake is minimal.
- 5. The lake-bottom sediment forms a thin seal that retards, but does not prevent, downward movement of water.

Conclusions not related to the objectives but that are consistent with those of other researchers are:

1. The deeper the piezometer is in the sediment, the greater the hydraulic differential.

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2. Water movement is primarily north and south of Austin Lake, with little east or west flow. This suggests that the divide is along an eastwest line that crosses Austin Lake and probably changes position during the year as the lake varies from recharging to discharging at particular points.

CHAPTER 6

RECOMMENDATIONS

Recommendations for future investigations include: 1. A study of the near shore hydrology. What are the effects of the layered sediments on local ground-water flow? Could these layers shed any light on historical development of the lake?

- 2. A study of conditions necessary for water movement into the lake. Does water move into the lake at all points or is it limited in extent near topographic highs or near the ground-water divide on the east and west sides?
- 3. A study of the water movement through the organic sediment filled basins. Are the basins impermeable? Do they act as sinks for dissolved materials?

REFERENCES

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- Allen, W. B., Miller, J. B., & Wood, W. B. Availability of water in Kalamazoo County, Southwest Michigan. U.S. Geological Survey Water Supply Paper, 1972,
- Born, S. M., Smith, S. A., & Stephenson, D. A. The hydrologic regime of glacial-terrain lakes, with management and planning applications. Madison: University of Wisconsin, 1974.
- Deutsch, M., Vanlier, K. E., & Giroux, P. R. Ground-water hydrology and glacial geology of the Kalamazoo area. Michigan. Michigan Geological Survey Progress Report, 1960, No. *2T.*
- Freeze, R. A., & Cherry, J. A. Groundwater. Englewood
Cliffs: Prentice Hall, Inc., 1979.
- Freeze, R. A., & Witherspoon, P. A. Theoretical analysis of regional ground-water flow: 2. Effect of watertable configuration and subsurface permeability variation. Water Resources Research, 1966, 3, 623-635.
- Hackbarth, D. A. Hydrogeology of the Little St. Germain
Lake basin, Vilas County, Wisconsin. Unpublished masrers thesis, University of Wisconsin, 1968.
- Jones .and Henry Engineers, Limited. Kalamazoo Metropolitan Planning Commission, Kalamazoo County, Michigan, water quality study. 1974.
- Lakes group set meeting. Kalamazoo Gazette, February 23,
1964, pp. 8.
- Lee, D. R. Septic tank nutrients in ground water entering Lake Sallie, Minnesota. Unpublished masters thesis, North Dakota University, 1972.
- Lee, D. R., & Cherry, J. A. A field exercise on ground-
water flow using seepage meters and mini-piezometers.
Journal of Geological Education, 1972, 27, 6-10.

Legrand, H. E. Patterns of contaminated zones of water in the ground. Water Resources Research, 1965, 1, 83-95.

- McBride, M. S., & Pfannkuch, H. 0. The distribution of seepage within lakebeds. Journal of Research of the
U.S. Geological Survey, 1975, 3, 505-512.
- Manson, P. W., Schwartz, G. M., & Allred, E. R. Some as-pects of the hydrology of ponds and small lakes. University of Minnesota Agriculture Experiment Station Bulletin, 1968, No. 257.
- Portage approves lakes level study. Kalamazoo Gazette, December 4, 1962. pp. 21.
- Rauch., V. Tri-lakes pipeline put in use. Kalamazoo Gazette, July 20, 1967, pp. 1, 2.
- Straw, W. T. Austin Lake: July 1977-January 1978. status report. Institute of Public Affairs, Western Michigan University, 1978.
- Toth, J. A theoretical analysis of ground-water flow in small drainage basins. Journal of Geophysical Re-
<u>search</u>, 1963, <u>68</u>, 4795-4812.
- Tri-lakes project under way. Kalamazoo Gazette, December 23, 1966, pp. 20.
- Wells may hold Long Lake level. Kalamazoo Gazette, September 21, 1963, pp. 1.
- Zeno, J. Austin Lake Project set. Kalamazoo Gazette, September 26, 1962, pp. **17**.

REFERENCE NOTES

1. Ryskamp, A. C. Personal communication, May, 1981.

APPENDIX A

PEARSON PRODUCT MOMENT CORRELATIONS OF ALL OBSERVATIONS COMBINED

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0.0453 COND

APPENDIX B

PEARSON PRODUCT MOMENT CORRELATIONS OF OBSERVATIONS CLASSIFIED ACCORDING TO VALUES ON INDEPENDENT VARIABLES— DISTANCE FROM SHORE AND PRESENCE OR ABSENCE OF A SEPTIC TANK

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APPENDIX C

PEARSON PRODUCT MOMENT CORRELATIONS OE OBSERVATIONS CLASSIFIED ACCORDING TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY

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APPENDIX D

ANALYSIS OF OBSERVATIONS CLASSIFIED ACCORDING TO VALUES ON BOTH INDEPENDENT VARIABLES USING A GENERAL LINEAR MODELS PROGRAM

STATISTICAL ANALYSIS SYSTEM 18:20 FRIDAY, OCTOBER 23, 19S1

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GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

NUHBER OF OBSERVATIONS IN DATA SET = 95

: VARIABLES IN EACH GROUP ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF HISSING VALUES.

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GENERAL LINEAR MODELS PROCEDURE

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GENERAL LINEAR MODELS PROCEDURE

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DEPENDENT VARIABLE: COND

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STATISTICAL ANALYSIS SYSTEM 16:20 FRIDAY, OCTOBER 21, 19B1

GENERAL LINEAR MODELS PROCEDURE

STATISTICAL ANALYSIS SYSTEM 18:20 FRIDAY, OCTOBER 23. 1981 10

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GENERAL LINEAR MODELS PROCEDURE

DEREMOENT VARIABLE: CUL

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APPENDIX E

ANALYSIS OP OBSERVATIONS CLASSIFIED ACCORDING TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY USING A GENERAL LINEAR MODELS PROGRAM

17:09 FRIDAY, OCTOBER 30, 1981 STATISTICAL ANALYSIS SYSTEM

GENERAL LINEAR MODELS PROCEDURE

VALUES CLASS LEVEL INFORMATION LEVELS **CLASS**

 $\ddot{}$ STNK NUMBER OF OBSERVATIONS IN DATA SET = 95

DEPENDENT VARIABLES aes anous

MOTE: VARIABLES IM EACH GROUP ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES.
STATISTICAL ANALYSIS SYSTEM 17:09 FRIDAY, OCTOBER 30. 19S1 \mathbf{r}

GENERAL LINEAR HODELS PROCEDURE

17:09 FRIDAY. OCTOBER 30. 1981 ATISTICAL ANALYSIS SYSTEM ST.

GENERAL LINEAR HOOELS PROCEDURE

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SENERAL LINEAR MODELS PROCEDURE

17:09 FRIDAY, OCTOBER 30, 1981 STATISTICAL ANALYSIS SYSTEM

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APPENDIX P

SCHEMATIC PLOTS FOR THE DEPENDENT VARIABLES

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APPENDIX G

EXAMINATION POR SKEWNESS USING THE UNIVARIATE PROCEDURE

11:54 SATURDAY, NOVEMBER 14, 1981 EXAMINATION OF DISTRIBUTIONS FOR SKENNESS UNIVARIATE

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11:54 SATURDAY, NOVEMBER 14, 1981

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EXAMINATION OF DISTRIBUTIONS FOR SKENNESS UNIVARIATE

TOMENTS VARIABLE=COND

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š Š 11:54 SATURDAY, NOVEMBER 14, 1981 EXAMINATION OF DISTRIBUTIONS FOR SKEWNESS UNIVARIATE

QUANTILES(DEF=4)

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Sun wars å **MONENTS**

VARIABLE=LNCOND

EXTREMES

BURGH

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Eneme
Province **RANGE**
Q3-Q1
RDDBE

0.908258
0.182322
5.16832

SSON/L % COUNT MISSING

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អ្នកទទួល 11:54 SATURDAY, NOVEMBER 14, 1981 11:54 SATURDAY, NOVEMBER 14, 1981 **EXTREMES EXTREMES** ក្នួននគតន
ខ្ម 998888 EXAMINATION OF DISTRIBUTIONS FOR SKENNESS EXAMINATION OF DISTRIBUTIONS FOR SKEMMESS **KNNNNN EXECUTE** NAMTILES(DEF=4) QUANTILES(DEF=4) SESSE 5°S 80666 99573 1,81093 UNIVARIATE **JNIVARIATE MISSING VALUE**
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APPENDIX H

PINAL ANALYSIS OP HYDRAULIC HEAD USING THE GENERAL LINEAR MODELS PROCEDURE

SECOND PRINT OUT-THAS DV AND DREADOES NOT HAVE DIST THIS SEEMS TO BE THE BEST MODEL FOR HH

13:23 SATUROAY, NOVEMBER 1

GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

NUMBER OF OBSERVATIONS IN DATA SET « 95

NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOWEVER, ONLY
OBSERVATIONS IN DATA SET CAN BE USED IN THIS ANALYSIS. \overline{a}

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14, 1981

 $\ddot{\cdot}$ 139.7498
HH MEAN
35632184

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APPENDIX I

FINAL ANALYSIS OF SPECIFIC CONDUCTIVITY,
CHLORIDE CONTENT, AND PHOSPHATE CONTENT
USING THE GENERAL LINEAR MODELS PROCEDURE

14:01 SATURDAY, NOVEMBER 14, 1981

FULL MOEL FOR LOG OF CONDUCTIVITY GENERAL LINEAR MODELS PROCEDURE CLASS LEVEL INFORMATION **CLASS**

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ទី VALUES ះ LEVELS

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NUMBER OF OBSERVATIONS IN DATA SET = 95

ã WOTE: ALL DEFEMOENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ADSENCE OF MISSING VALUES, NOWEVER, ONLY
// COSERVATIONS IN DATA SET CAN BE USED IN THIS ANALYSIS,

108

FULL MODEL FOR LOG OF CONDUCTIVITY

14:01 SATURDAY, NOVEMBER 14, 1981

GENERAL LINEAR MODELS PROCEDURE

LEAST SQUARES MEANS

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0.1323
0.1323
0.000

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED. 2 0.1455 0.1455 0.1679
2 0.1455 0.1959 0.7959 등등등
- 000
- 000 0.07717500
0.04440764
0.05042015 5.65658257
5.76831000
5.75648829 -0.5

PROB > |T| HO:
LSHEAN1=LSHEAN2

РЯОВ > |I|
НО:LSMEAN=0

STD ERR
LSMEAN

LACOND
LSHEAM 5.7408800
5.71336591 LACOND 1,67771814

DIST

0.5173

PROB > 171
LSHEAN1=LSHEAN2

PROB > 111
HO:LSMEAN=0 0.0001

STD ERR
LSMEAN 0.04797368
0.05187684

STHK

 0.0277

 $\frac{800}{3000}$

0.05354731
0.04665309

ć

109

REDUCED MODEL FOR LOG OF CONDUCTIVITY--HH AND STAK ONLY
MODEL INCLUDES TEST FOR INTERACTION BETWEEN HH AND STAK

14:01 SATURDAY, NOVEMBER 14, 1981

GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

VALUES

LEVELS **CLASS**

 $\frac{1}{\alpha}$ $\ddot{}$ **STINK** NUMBER OF OBSERVATIONS IN DATA SET = 95

š NOTE: ALL DEPENDEMT VARIABLES ARE COMSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSIMO VALUES. NONEVER, ONLY
DSSERVATIONS IN DATA SET CAM BE USED IN THIS ANALYSIS.

2.5068

0.107320 R-SQUARE

3.29

F VALUE

 $\frac{1}{2}$ 0.0246 STD DEV 0.14973840

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ن LNCOND MEAN 5.78858207 $\frac{1}{2}$ ing
Consta

ន្លុះទុ **F VALUE**

 $0.0208003
0.13468570
0.02309510$ TYPE IV SS

SONGLOC

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PROB > |T| HO:
LSHEAN1=LSHEAN2 0.0770 PROB > |T|
HO:LSMEAN=0 0.0007
0.0007 LEAST SQUARES HEANS STD ERR
LISMEAN 0.02100802
0.02734543 LHCOND 5.76069992
5.82245504 **STINK**

14:01 SATURDAY, NOVEMBER 14, 1981

FULL HODEL FOR LOG OF CHLORIDE GENERAL LINEAR MODELS PROCEDURE CLASS LEVEL INFORMATION VALUES $\ddot{}$ LEVELS CLASS ಕ

1 5 61 62 63 64 $\ddot{}$ SO₁ **bist** STINK NUMBER OF OBSERVATIONS IN DATA SET = 95

å NOTE: ALL DEFENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES, HOMEVER, OMLY
NOTE: OSSERVATIONS IN DATA SET CON BE USED IN THIS ANALYSIS.

JENERAL LINEAR MODELS PROCEDURE LEAST SQUARES HEANS

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.

PROB > ITI HO: LSMEAN(1)=LSMEAN(J)
I/J > 1

PROB > |T|
HO LSMEAN=0

STD ERR
LSMEAN 0.06629866
0.09237738
0.06528690
0.00819223
0.008921968

 $\frac{1}{8}$ --5855

1.51135602
1.53637960
1.47289426
1.40427498
1.39084214 **LARGE**

FULL HODEL FOR LOG OF CHLORIDE

14:01 SATURDAY, NOVEMBER 14, 1981

РАОВ > 111 НО: LSMEAN(1)=LSMEAN(J)

PROB > |T|
HO:LSMEAN=0

2 0.6274 0.6274 0.5450
2 0.6274 0.9839 0.9839

들들
이 이 이 이
이 이 이

0.00249541
1400470047
1400470047 STD ERR
Lishean

3.44599750
3.47311118
3.47438788 LSHEMI

d $\ddot{}$ 112

ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD ONLY PROBABILITIES PROTECTION LEVEL, **NOTE: TO ENSURE OVERALL**

PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED CONPARISONS SHOULD BE USED.

LYI
ISBN > III HO: TRNEVA(I)=TRNEVA(1)

PROB > ITI
HO:LSHEAN=0

STD ERR
LISHEAN 0.06340807
0.03650618
0.04105315

PHIPL

HOTE: TO ENSURE OVERALL

3.50383082
3.40751445
5.40215129 **LHEAT**

0.7123
0.0467

0.1268 0.0467 $\begin{bmatrix} 1 & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 0 \\$

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0.9899
0.3637

0.8506

compared 0.7319

0.2226
0.3942

0.8643

REDUCED MODEL FOR LOG OF CHLORIDES--HH AND STHK ONLY
MODEL INCLUDES TEST FOR INTERACTION BETWEEN HH AND STHK

GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

VALUES LEVELS **CLASS**

t \sim STNK

NUMBER OF OBSERVATIONS IN DATA SET = 95

ä NOTE ALL DEFROURNT VARIABLES ARE CONSISTENT N'IN RESPECT TO THE PRESENCE OR AGSENCE OF MISSING VALUES. NONEVER, ONLY
- OSSERVATIONS IN DATA SET CAN BE USED IN THIS ANALYSIS.

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0.000

0.01824054
0.02330795

3.39658611
3.44178400

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14:01 SATURDAY, NOVEMBER 14, 1981

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OF OBSERVATIONS IN DATA SET » 95

NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOWEVER, ONLY 82
OBSERVATIONS IN DATA SET CAN BE USED IN THIS ANALYSIS.

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14:01 SATURDAY, NOVEMBER 14, 1981

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GEHERAL LINEAR MODELS PROCEDURE

FULL MODEL FOR LOG OF PHOSPHATES

LEAST SQUARES MEANS

 $\begin{array}{cccc} 0.9704 & 2 & 0.4146 & 0.068 & 0.0168 \ 0.0001 & 3 & 0.0240 & 0.0168 & \end{array}$ 0.00369804 0.09928908
0.24165080 0.05451940 . .

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANHED COMPARISONS SNOULD BE USED. PROB > 111 HO: LSMEAN(1)=LSMEAN(J)
1/J
1 PROB > 1T1
HO:LSNEAN=0 STD ERR
LSMEAN **HEAR**
SHEAR soi

0.6439 0.5054 0.8923
0.6366
0.3373 ime
Sine 0.2905
 0.6343 0.5619 0.8368 ន្លិន្តិ ទីចំ 0.07563328
0.04592133 1015998 53 0.10737810
0.14715864 01278890 ..5833 NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.

. 0.0112 3 0.0583 0.9327 0.06370308 0.16632900

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.

REDUCED HODEL FOR LOG OF PHOSPHATES-HH AND STHK ONLY
MODEL INCLUDES TEST FOR INTERACTION BETHERN HH AND STHK

GENERAL LINEAR MODELS PROCEDURE

14:01 SATURDAY, NOVEMBER 14, 1981

CLASS LEVEL INFORMATION

VALUES LEVELS **CLASS**

 $\frac{1}{2}$ \sim **STINK**

UNDER OF OBSERVATIONS IN DATA SET = 95

á NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT NITH RESPECT TO THE PRESENCE OR ADSENCE OF MISSING VALUES, RONEVER, ONLY
DESERVATIONS IN DATA SET CAN BE USED IN THIS ANALYSIS. š. $2.5.2$ L. 522

PROB > ITI
LSHEAN1=LSHEAN2 0.1883

PROB > |T|
HO:LSMEAN=O 0.0001

STD ERR
LSHEAN 0.02806086
0.03613032

STAK

0.22903203
0.16913695 HH LHP
HLJ

BIBLIOGRAPHY

- Allen, W. B., Miller, J. B., & Wood, W. B. Availability of water in Kalamazoo Comity, Southwest Michigan. U.S. Geological Survey Water Supply Paper, 1972. $No. 1973.$
- Allred, E. R., Manson, P. W., Schwartz, G. M., Golany, P, & Reinke, J. W. Continuation of studies on the hydro-logy of ponds and small lakes. Minnesota University of Agriculture Experiment Station Bulletin, 1971,
No. 274.
- Born, S. M., Smith, S. A., & Stephenson, D. A. The hydrologic regime of glacial-terrain lakes, with management and planning applications. Madison: University of Wisconsin, 1974.
- Crosby, J. W., Johnstone, D. L., Drake, C. H., & Eenton, R. L. Migration of pollutants in a glacial outwash
environment. Water Resources Research, 1968, 4, 1095-1113.

 \mathcal{I}

- Deutsch, M. Ground-water contamination and .legal controls in Michigan. U.S. Geological Survey Water Supply Paper, 1963, No. 1691.
- Deutsch, M., Vanlier, K. E., & Giroux, P. R. Ground-water hydrology and glacial geology of the Kalamazoo area, Michigan. Michigan Geological Survey Progress Report, 1960, No. 23.
- Eischer, F. E. Fundamental statistical concepts. San Francisco: Canfield Press, 1973-
- Freeze, R. A., & Cherry, J. A. Groundwater. Englewood Cliffs: Prentice Hall, Inc., 1979.
- Freeze, R. A., & Witherspoon, P. A. Theoretical analysis of regional ground-water flow: 1. Analytical and numerical solutions to the mathematical model. Water Resources Research, 1966, 2, 641-656.
- Freeze, R. A., & Witherspoon, P. A. Theoretical analysis of regional ground-water flow: 2. Effect of watertable configuration and subsurface permeability variation. Water Resources Research, 1966, 3, 623-635.

- Hach Chemical Company. Hach DR colorimeter methods manual. 1977, 13th edition.
- Hackbarth, D. A. Hydrogeology of the Little St. Germain
Lake basin, Vilas County, Wisconsin. Unpublished masters thesis, University of Wisconsin, 1968.
- Jones and Henry Engineers., Limited. Kalamazoo Metro-.. politan Planning Commission, Kalamazoo County, Mich-igan, water quality study. 19?4.
- Lakes group set meeting. Kalamazoo Gazette. Februarv 23. $1964, pp.8.$
- Lee, D. R. Septic tank nutrients in ground water entering Lake Sallie. Minnesota. Unpublished masters thesis. University of North Dakota, 1972.
- Lee, D. R. A device for measuring seepage flux in lakes and estuaries. Limnology and Oceanography, 1977, 22, 140-147. 140–147.
- Lee, D. R., & Cherry, J. A. A field exercise on groundwater flow using seepage meters and mini-piezometers. Journal of Geological Education, 1972, 27, 6-10.
- Legrand, H. E. Patterns of contaminated zones of water in the ground. Water Resources Research, 1965, 1, 83-95. 83-95* ----------------------
- McBride, M. S., & Pfannkuch, H. O. The distribution of seepage within lakebeds. Journal of Research of the
U.S. Geological Survey, 1975, 2, 505-512.
- Manson, P. W., Schwartz, G. M., & Allred, E. R. Some aspects of the hydrology of ponds and small lakes. University of Minnesota Agriculture Experiment Station Bulletin, 1968, No. 257.
- Portage approves lakes level study. Kalamazoo Gazette, Decomber 4, 1962, pp. 21.
- Rauch, V. Tri-lakes pipeline put in use. Kalamazoo Gazette, July 20, 1967, pp- 1, 2.
- Ryskamp, A. C. Personal communication, May, 1981.
- SAS Institute Inc. SAS user's guide. 1979 edition.
- Sayre, A. N. Groundwater. Scientific American, November 1950, 183, No. 5, 14-19.
- Straw, W. T. Austin Lake: July 1977-January 1978. A
status report. Institute of Public Affairs, Western Michigan university, 1978.
- Toth, J. A theoretical analysis of ground-water flow in small drainage basins. Journal of Geophysical Re-
<u>search</u>, 1963, <u>6</u>8, 4795-4812.
- Tri-lakes project under way. Kalamazoo Gazette, December 23, 1966, pp. 20.
- Wells may hold Long Lake level. Kalamazoo Gazette, Septem-ber 21, 1963, pp. 1.
- Winter, T. C. Numerical simulation of steady state threedimensional ground-water flow near lakes. Water Resources Research, 1978, 14, 245-254.
- Zeno, J. Austin Lake project set. Kalamazoo Gazette, Sep-tember 26, 1962, pp. **17**.