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GROUND-WATER FLOW ALONG THE TOTAL SHORELINE OF AUSTIN LAKE AND ITS ENVIRONMENTAL CONTRIBUTION TO POLLUTION OF THE LAKE

by

William Thomas Williams

A Dissertation Submitted to the Faculty of The Graduate College in partial fulfillment of the requirements for the Degree of Doctor of Philosophy Science Education

Western Michigan University Kalamazoo, Michigan April 1982

GROUND-WATER FLOW ALONG THE TOTAL SHORELINE OF AUSTIN LAKE AND ITS ENVIRONMENTAL CONTRIBUTION TO POLLUTION OF THE LAKE

William Thomas Williams, Ph.D. Western Michigan University, 1982

The vertical direction of ground-water flow in Austin Lake was determined by placing minipiezometers along the entire shoreline at fifty randomly selected sites. Groundwater samples were taken at each site and the specific conductivity of the water samples was determined along with the chloride and phosphate content. The data were analyzed to determine if there were any relationships among water movement, chemical content, and the presence or absence of an onshore septic tank. The minipiezometer data show that the primary direction of water movement through the sediment-water interface is downward but may become reversed at localized regions as conditions vary. The negative hydraulic differential is greater at the north and south ends of the lake than on the east and west sides. This observation makes a stronger case for reports by other researchers that Austin Lake is on the ground-water divide and that ground water flows to the north and to the south from the lake. Computer analyses of all data show relationships among specific conductivity. chloride content, and the presence of a septic tank. Offshore from septic tanks, the specific conductivity and chloride content

vary more and have higher average concentrations than offshore from areas without a septic tank, although visual inspection of the data shows pollution to be minimal. Phosphate content failed to show relationships with septic tanks, but tended to be higher on the west side of the lake--an area bounded by outwash and containing older functional and non-functional septic tanks.

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William Thomas Williams

ii

TABLE OF CONTENTS

And the second second second

ACKNO	LEDG:	EMENI	s.	•••	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	٠	ii
LIST	OF TA	BLES	•			•	•		•		•		•	•	•		•	•		•	vi
LIST	OF FI	GURES	5.		•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	vii
Chapt	er																				
1.	INT	RODUC	TIC	ON.	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	•	1
	Geo	graph	nica	al	Loc	cat	tic	n	•		•	•	•	•	•	•	•	•	•	•	· 1
	Geo	logy	•				•	•	•		•	•	•		•			•	•		1
		rces	of	Wa	tei	r]	Ent	er	rir	ıg	ar	nđ	Le	a,	rir	ıg	Αı	ist	;ir	1	3
		ation		fS	ew.	- -	т.4	ne		Ī	Ī	ļ	ļ								10
		tory		1~~		-	101			•	·	·	•	·	•	•	•	·	·	•	10
		v		•••	•	·	•	•	•	•	•	•	•	•	•	•	•	•	•	•	
	Ass	essme	ent	of	L	ak	eΕ	۲n	riı	or	nme	ent	ts	•	•	•	•	•	•	•	12
	Sig	nific	an	ce	of	S	tuċ	ly	•	•	•	•	•	•	•	•	•	•	•	•	13
2.	REV	IEW C	OF 3	PER	TI	νE	NΤ	Ľ	TI	ER.	ΥT	JR]	Ξ.	•	•	•		•	•	•	15
	Gro	und-v	vat	er	Flo	ow	•		•	•	•	•	•	•	•	•	•	•	•	•	15
	Gro	und-v	vat	er	Qua	al	itʒ	<i>.</i>	•	•	•	•	•	•	•	•	•	•	•	•	25
3.	PRO	CEDUF	RE :	FOF	2 C(OL:	LEC). T	INC	3 1	DA'	ΓA	•	•	•	•	•	•	•	•	28
	Det	ermir	nat	ior	1 03	f	Hyć	ira	au.	Lio	c F	Iea	ad	•	•	•	•	•	•	•	28
	Wat	er Sa	amp	lir	ıg.	•	•	•	•	•	•	•	•	•	•		•	•	•	•	32
	Che	mical	L A	nal	.ys	is	•	•	•	•	•	•	•	•		•	•	•	•	•	32
		ortar																			33

iii

TABLE OF CONTENTS (continued)

4.	ANALYSIS OF DATA	35
	Table Descriptions	35
	Variable Descriptions	47
	-	
	Analytical Procedure	48
	Computer Results	50
	Hydraulic Head	50
	Conductivity	54
	Chlorides	56
	Phosphates	57
5.	DISCUSSION AND CONCLUSIONS	58
	Ground-water Flow	58
	Pollution	63
	Summary of Conclusions	69
6.	RECOMMENDATIONS	72
REFEREN	NCES	73
REFEREN	NCE NOTES	75
APPENDI	IX A	76
PEA	ARSON FRODUCT MOMENT CORRELATIONS OF ALL	
	OBSERVATIONS COMBINED	76
APPENDI	IX B	78
PEA	ARSON PRODUCT MOMENT CORRELATIONS OF OBSERVA-	
	TIONS CLASSIFIED ACCORDING TO VALUES ON IN- DEPENDENT VARIABLES-DISTANCE FROM SHORE	
	AND PRESENCE OR ABSENCE OF A SEPTIC TANK	78

TABLE OF CONTENTS (continued)

APPENDIX C	83
PEARSON PRODUCT MOMENT CORRELATIONS OF OBSERVA- TIONS CLASSIFIED ACCORDING TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY.	83
	-
APPENDIX D	86
ANALYSIS OF OBSERVATIONS CLASSIFIED ACCORDING TO VALUES ON BOTH INDEPENDENT VARIABLES	
	36
APPENDIX E	90
ANALYSIS OF OBSERVATIONS CLASSIFIED ACCORDING	
TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY USING A GENERAL LINEAR MODELS PROGRAM.	90
APPENDIX F	94
SCHEMATIC PLOTS FOR THE DEPENDENT VARIABLES 9	94
APPENDIX G	99
EXAMINATION FOR SKEWNESS USING THE UNIVARIATE	
PROCEDURE	99
APPENDIX H	04
FINAL ANALYSIS OF HYDRAULIC HEAD USING THE	
GENERAL LINEAR MODELS PROCEDURE 10	04
APPENDIX I	27
FINAL ANALYSIS OF SPECIFIC CONDUCTIVITY,	
CHLORIDE CONTENT, AND PHOSPHATE CONTENT USING THE GENERAL LINEAR MODELS PROGRAM 10	27
BIBLIOGRAPHI	17

v

LIST OF TABLES

Table 1:	Physical data for each study site	36
Table 2:	Water analysis results for each study site	40
Table 3:	Factors considered in the statistical analysis of Austin Lake's ground-water characteristics	44

vi

LIST OF FIGURES

Figure	1:	Location of study site in Kalamazoo County, Michigan 2
Figure	2:	Cross sections of Austin Lake 4
Figure	3:	Easal sediment in Austin Lake 5
Figure	4:	Drainage basin of Austin Lake 7
Figure	5:	Water sources and losses of Austin Lake
Figure	6:	The primary drainage basins of Kala- mazoo County, Michigan 9
Figure	7: ⁻	Simplified flow pattern of ground water showing ground-water divides 16
Figure	8:	Flow lines and equipotential lines 18
Figure	9:	Effect of topography on regional ground-water flow patterns 19
Figure	10:	Configuration of ground-water flow systems around lakes
Figure	11:	Ground-water divides as shown by Deutsch et al. and Allen et al 24
Figure	12:	Lateral spread of pollutants in the zone of saturation
Figure	13:	Approximate locations of sample sites on Austin Lake
Figure	14:	Sampling equipment
Figure	15:	Hydraulic head at each sampling site around Austin Lake
Figure	16:	Possible ground-water flow patterns at Austin Lake 61
Figure	17:	Average hydraulic head readings for the different types of glacial mate- rial

vii

LIST OF FIGURES (continued)

Figure	18:	As the hydraulic head becomes more positive the conductivity increases 65
Figure	19:	A comparison of the specific conduc- tivity with the presence or absence of a septic tank
Figure	20:	A comparison of the chloride concen- trations with the presence or absence of a septic tank
Figure	21:	Phosphate concentration averages for the different types of glacial mate- rial

viii

CHAPTER 1

INTRODUCTION

Geographical Location

Austin Lake is located in the southeastern corner of Portage Township, Kalamazoo County, Michigan (Figure 1). The lake encompasses parts of sections 23, 24, 25, 26, 35, and 36, T. 3 S., R. 11 W.

Geology

Austin Lake lies in a region glaciated by a succession of at least four major continental ice sheets. The main topographical features of the area were formed during the recession of the last ice sheet. Austin Lake is a "kettle" lake, occupying a depression formed from the melting of an isolated block, or blocks, of ice left during the retreats and readvances of the Lake Michigan lobe of the glacier during the last Ice Age--Wisconsinan time (Deutsch, Vanlier, & Giroux, 1960). The three-dimensional shape of the basins suggests that the lake was formed by the melting of three blocks of ice that were partly connected (Straw, 1978). Organic sediments have filled the lake basins until at the present time the maximal surface water depth is less than 10 feet (3.0 m). The

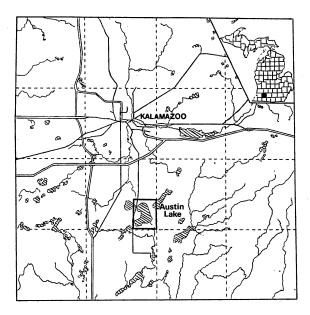


Figure 1. Location of study site in Kalamazoo County, Michigan.

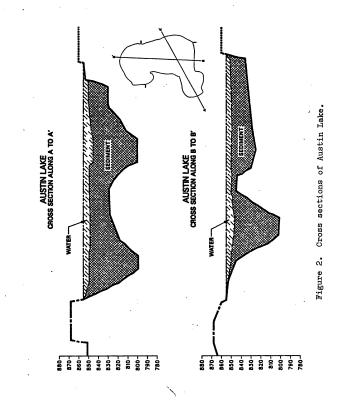
basins formed by the three blocks of ice range in depth from 25 feet (7.6 m) to 50 feet (15.2 m) with the deepest being in the southwestern portion of the lake (Figure 2). Below the organic material in the basins, the sediment varies from coarse gravel to clay (Figure 3). Along most of the shoreline, the sediments have been modified by waves so that sand and gravel predominate.

Surface glacial deposits surrounding Austin Lake are of three types--morainal, outwash, and channel deposits. Adjacent to the lake on the northeast side are morainal deposits that trend northeastward. These morainal deposits were deposited directly from glacial ice with running water playing a minimal role in deposition. The west and south sides consist of outwash--material that was deposited from meltwater streams issuing from glacial ice. On the east side of the lake, trending southeastward, and from the north part of the lake northwestward are channel deposits. These are similar in origin to the outwash deposits, but have been reworked and covered with finer materials (Deutsch et al., 1960).

Sources of Water Entering and Leaving Austin Lake

Austin Lake covers an area of 1050 acres (2593.5 ha), but drains a land area that encompasses only 340

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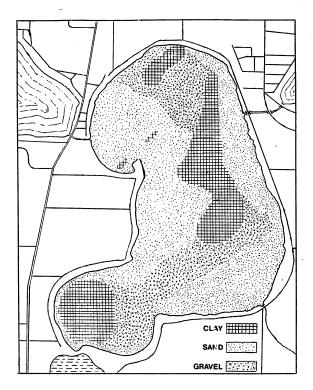


Figure 3. Basal sediment in Austin Lake (Straw, 1978).

acres (839.8 ha) (Straw, 1978). The effective drainage basin is considerably smaller because the permeability of the surrounding soil is high (Figure 4). Some surface water enters the lake from storm sewers, but otherwise surface water from the immediate drainage basin contributes little to the water budget. Some small ephemeral streams are present in areas of moderate to steep slopes, but most such drains do not contribute water unless rainfall is heavy. Other sources of water include canals from Long Lake and West Lake, direct precipitation, the Upjohn recharge pond, surface condensation, and ground-water discharge.

An outlet at the southeastern end of the lake leads to Gourdneck Creek. Water also seeps through the bottom to become part of the ground-water flow. Some water is lost through evaporation from the free water surface and through irrigation (Figure 5).

Austin Lake, with the other lakes in the complex, are all considered to be part of the St. Joseph River basin (Figure 6). Yet, Allen, Miller, and Wood (1972) placed the ground-water divide (major deep ground-water flow) on an east-west line through the center of Austin Lake with ground-water movement to the north and south of the lake (Figure 11).

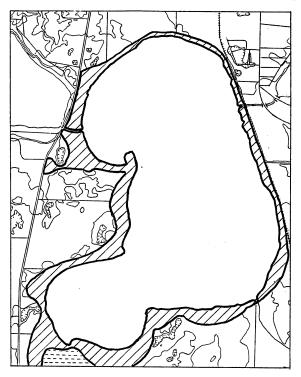


Figure 4. Drainage basin of Austin Lake (Straw, 1978).

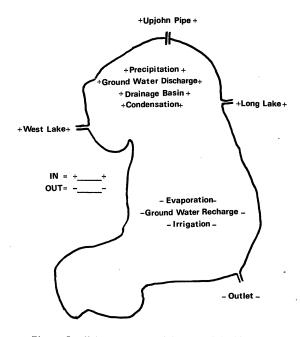


Figure 5. Water sources and losses of Austin Lake.

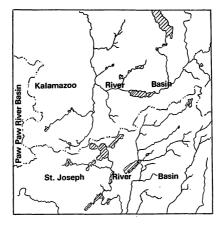


Figure 6. The primary drainage basins of Kalamazoo County, Michigan (Allen et al., 1972).

Location of Sewer Lines

The dwellings around Austin Lake are served by sewer lines with the exception of those in the southwestern corner. Because sewer lines surrounding the lake are of recent origin many residents still use septic tanks.

History

Natural accumulations of organic sediments have partly filled the lake since the retreat of the last glacier. Radiocarbon dating of the organic sediments in certain parts of the lake indicate that the rate of accumulation of these materials was about 1 inch (2.54 cm) every 20 years for the period 8000 B. P. (before present) to 3000 B. P., 1 inch (2.54 cm) every 12 years for the period 3000 B. P. to 1700 B. P., and 1 inch (2.54 cm) every 7 years since 1700 B. P. It is likely that sedimentation may now be taking place at the rate of 1 inch (2.54 cm) every 2 years (Straw, 1978).

In 1962 an attempt was made to pump out the organic sediment near the north end of the lake. The objectives of the project were to determine (1) how much solid material could be pumped out in one hour using a four inch pipe, (2) if the hole formed by the pumping of muck would be filled in by the surrounding material, and (3) at what rate the hole would fill (Zeno, Kalamazoo

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<u>Gazette</u>, September 26, 1962). A holding area was constructed and filled with the pumped muck. The holding area was permeable so the water could drain out leaving the solids behind. After being filled, the weather turned cold, freezing and damaging the equipment. No further action was taken (Ryskamp, Note 1).

Also in 1962, concern over the lake level led to a plan for a ponding project to hold water to be released when the level of the lakes dropped (<u>Kalamazoo Gazette</u>, December 4, 1962). In 1963 attention was directed towards pumping well water into the lakes to maintain their levels (<u>Kalamazoo Gazette</u>, Swptember 21, 1963). In 1964 a proposal was presented to divert water from the Upjohn pond to Austin Lake (<u>Kalamazoo Gazette</u>, February 23, 1964). This became known as the Tri-Lakes Upjohn Project and, after a series of legal battles, construction finally began in December of 1966 (<u>Kalamazoo Gazette</u>, December 23, 1966). On July 20, 1967, (<u>Kalamazoo Ga</u>-<u>zette</u>) the pipeline construction was completed and five million gallons of water a day began flowing into Austin Lake.

The pipeline was instrumental in solving the low water-level problem, but because the lake is shallow, sunlight is able to penetrate to the bottom stimulating the growth of rooted aquatic plants. Changes in the use

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of the lake from predominantly fishing to fishing, boating, and water skiing caused residents to view the growth of weeds with concern. Various methods have been proposed to alleviate this problem. These methods include dredging, aeration, chemical treatment, and harwesting, but costs and uncertainity of the results have prevented any action. More basic research to reduce the degree of uncertainity needs to be done before major steps can be taken to improve the lake basin.

Assessment of Lake Environments

This paper is a result of an investigation into the movement and quality of ground water as it relates to Austin Lake. The purpose of the investigation was to locate areas of inflow and outflow near the shoreline of Austin Lake and to determine if the lake is being polluted from ground-water sources.

The main hydrogeologic factors that are prerequisites for assessing lake environments are: (1) regime dominance (the relative magnitude of ground water in the water budget of a lake); (2) system efficiency (a term describing the rate aspects of surface- and ground-water movement through a lake system); and (3) position within a ground-water flow system (Born, Smith, & Stephenson, 1974). This paper is concerned with this last factor

along with the water quality of the ground water that enters the lake.

The best information for determining the position of a lake in a ground-water flow system is the groundwater potential distribution beneath the lake bottom, and the vertical hydraulic gradient is the ideal criterion. If the vertical gradient is positive (potential increasing downward) the lake is in a ground-water discharge region. If the vertical gradient is negative (potential decreasing downward) the lake is in a groundwater recharge region. If the vertical gradient is zero or negative at one end and positive at the other, the lake is in a region of lateral flow. Measurement of this factor must be done onsite with piezometers (Born et al., 1974).

1

Significance of Study

This study is significant for three reasons: (1) little research has been done on the ground-water flow beneath lakes in Michigan and none on Austin Lake with the exception of generalized deep ground-water flow patterns. The ground-water flow into and out of the lake has not been determined in detail by use of piezometers; (2) the lakeshore is highly urbanized and there is the possibility of substantial pollution; and (3) a detailed

knowledge of the hydrology of Austin Lake must be gained before an adequate plan can be developed for the restoration of the lake.

CHAPTER 2

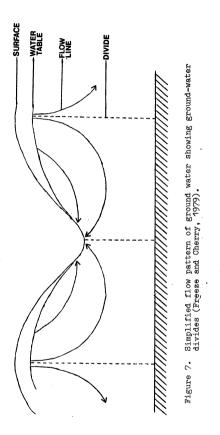
REVIEW OF PERTINENT LITERATURE

Ground-water flow and its association with lakes has only recently become the object of detailed research studies. Because of this, information on the subject is limited. Research on the quality of ground water that enters lakes is virtually non-existent. This chapter reviews research and writings of ground-water flow as it relates to lakes.

Ground-water Flow

Freeze and Cherry (1979) described ground-water movement in homogenous and isotropic materials as a flow from highlands towards valleys. In this situation the water table forms a subdued replica of the topography. The symmetry of the system creates vertical boundaries beneath the valleys and ridges across which there is no flow. These "ground-water divides" usually coincide exactly with surface-water divides, and their orientation is, for all practical purposes, vertical (Figure 7). Freeze and Weatherspoon (1966) showed that in more complex topographic and hydrogeologic environments groundwater movement becomes more complicated.

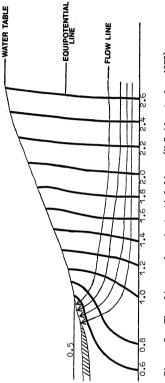
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Toth (1963) and Born et al. (1974) described groundwater flow as ground-water movement in response to gravity from regions of high to low potential along flow lines that describe the path of an individual water particle as it moves through the soil. A set of flow lines oriented so that any two flow lines adjacent at one point in the system remain adjacent throughout the system make up a ground-water flow system. All the flow lines in a ground-water flow system can be intersected by an uninterrupted surface across which flow takes place in one direction only. Freeze and Cherry (1979) stated that flow lines deliver ground water from recharge regions to discharge regions (Figure 8).

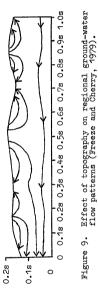
Freeze and Cherry (1979) described ground-water flow beneath a hilly topography. In this situation there are numerous subsystems within a major flow system. Water that enters the flow system in a given recharge region may be discharged in the nearest topographic low-(local ground-water flow) or it may be transmitted to the regional discharge region in the bottom of the major valley. Toth (1963) showed that as the depth to lateral extent of the entire system becomes smaller and as the amplitude of the hummocks become larger, the local systems are more likely to reach the basal boundary, creating a series of small independent cells (Figure 9).

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Flow lines and equipotential lines (McBride et al., 1975). Figure 8.

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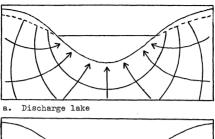


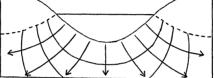
Freeze and Cherry (1979) listed the effects geological heterogeneity can have on regional ground-water flow. Heterogeneity can affect the interrelationship among local and regional ground-water flow, the surficial pattern of recharge and discharge regions, and the quantities of flow that are discharged through the systems.

Born et al (1974) discussed lakes and their association with ground water. Lakes situated in ground-water recharge regions (recharge lakes) contribute to the ground water throughout the entire lake bottom. Lakes in ground-water discharge regions (discharge lakes) gain ground water throughout the entire lake bottom. In regions of lateral ground-water flow, lakes lose to the ground water on one side and gain ground water on the other side. Such lakes are known as flow-through lakes (Figure 10). Lakes may be associated with the regional ground-water flow, local ground-water flow, or both. It is possible for lakes to intersect both a shallow and a deep flow system with the result that ground-water movement near the lake margins may be opposite to the directions of movement near the lake center.

Manson, Schwartz, and Allred (1968) indicated that "muck" is highly impermeable and therefore, where it exists, prevents water from moving in and out of a lake. They stressed the importance of bottom materials in preventing or permitting ground water to enter or to leave

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b. Recharge lake

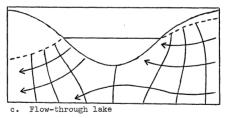


Figure 10. Configuration of ground-water flow systems around lakes (Born et al., 1974)

lakes. They also speculated that glacial lakes--not intersecting the water table--upon initial formation did not hold water. Finer materials carried in by sheet and gully erosion sealed the bottom of the kettles, thereby permitting the lakes to hold water. They presented data to support the view that the permeability of the soil and drift adjacent to the pond or lake are usually much greater than that of the bottom. Thus, when the water level rises to flood the normal basin, infiltration is rapid and the level tends to decline rather promptly to the usual level. Their investigation of ponds and small lakes in relatively deep depressions showed that inflow underground was largely limited to wet periods--and even then was not prevalent at all sites.

McBride and Pfannkuch (1975) and Lee (1972) showed that seepage into or out of lakes tends to be concentrated near the shore and decreases with increasing distance from the shore. In many places the rate of decrease is exponential. Their model indicated that little of the ground-water flow passes beneath the lake. Instead, most of it is diverted upward as it approaches the lake, passes through the lake, moves downward, and eventually comes back to the horizontal. McBride and Pfannkuch (1975) showed that this holds true only where the diameter of the lake is at least roughly comparable to the thickness of the underlying permeable materials.

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Where ponds have diameters much less than the thickness of the sections, seepage occurs almost uniformly across the bottoms of the depressions.

McBride and Pfannkuch (1975) in citing Lee's (1972) study of Lake Sallie, Minnesota, noted that in the 2 km wide lake, one-half of the discharge occurs within 17 m of shore, 90 percent within 60 m and 99 percent within 120 m. The edge of the fine-grained sediments in the center of Lake Sallie is 300 m from shore. This and similar research done by Hackbarth (1968) in the Little St. Germain Lake Basin in Wisconsin show the importance of the geometry of the ground-water system and the relatively small role that lake sediments have in controlling the distribution of seepage in lakes.

Deutsch et al. (1960) and Allen et al. (1972), using water level data from observation wells, lakes, and streams, determined the direction of the regional groundwater flow in the Austin Lake complex. This complex includes Austin, West, Long, Sugarloaf, and Gourdneck lakes. These data indicated that the ground-water movement was both north and south of an east-west line crossing the approximate center of Austin Lake (Figure 11).

Allen et al. (1972) devised a water budget for the complex from October 1, 1965 to March 30, 1966 and again from April 1, 1966 to September 30, 1966. For their first budget, they cited a subsurface inflow of 73 acre-feet

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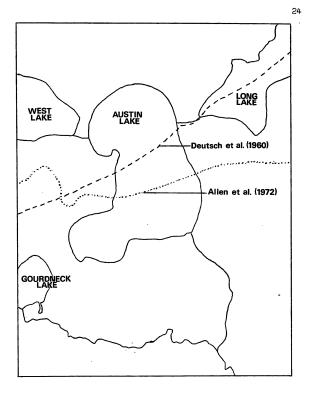


Figure 11. Ground-water divides as shown by Deutsch et al. (1960) and Allen et al. (1972).

and a subsurface outflow of 1,070 acre-feet. For the second budget, the subsurface inflow is 146 acre-feet and the subsurface outflow is 1,010 acre-feet. They said the subsurface outflow returns to the surface at Gourdneck Creek, a short distance to the south of Austin Lake and flowing east. These figures indicate that little ground water enters Austin Lake as compared with the total flow.

Ground-water Quality

Legrand (1965) described contaminant movement from waste sites into the subsurface water circulation. This movement results in contaminated zones. or enclaves. in the zone of saturation. The contaminants tend to be entrained in ground-water flow and also tend to attenuate to varying degrees by dilution in water, decay with time, or some other 'die away' mechanism, and adsorption on earth materials. Once a contaminated zone is approximately stable, an increase in concentration may cause it to remain about the same size under some conditions and enlarge in others according to combined attenuation effects. Where attenuation occurs only by dilution the contaminated zone will become enlarged with increased concentration, and even if attenuated only slightly by dilution may become greatly elongated in the direction of ground-water flow. Where attenuation occurs through

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decay or adsorption or through both mechanisms, the contaminated zone may not change appreciably, even if the concentration of contaminants increases. Where the contaminants sink into the ground, the vertical direction predominates in the zone of aeration that lies above the water table. When they reach the zone of saturation contaminants tend to spread laterally (Figure 12).

Lee (1972), in his study of lakes in a glacial terrain, found that septic tank effluent from lakeside houses entered the lake within 9 meters of the shoreline in a fan-shaped pattern. His study showed that 40 percent of the nitrogen entered the lake although the phosphorus was fixed in the soil.

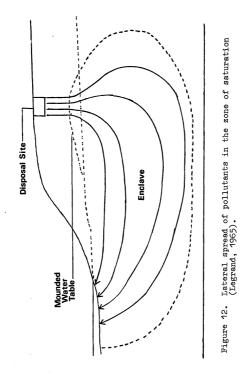
Previous studies of the water quality of Austin Lake indicate low concentrations of dissolved nutrients. Jones and Henry Engineers, Limited (1972) showed that chloride concentrations on the western edge of the lake averaged 36 mg/l and 20 mg/l at other sampling points. Orthophosphates ranged from 0 to 0.45 mg/l.

Allen et al. (1972) listed chloride concentrations between 1 and 6 mg/l.

Chemical analyses by the author during a period from September to December, 1977, show low levels of phosphorus.

Information on the quality of ground water discharged into Austin Lake is not available.

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CHAPTER 3

PROCEDURE FOR COLLECTING DATA

The purpose of this investigation was twofold. The first objective was to determine the direction of water flow through the sediment-water interface along the total shoreline of Austin Lake. The second objective was to determine if this water was contributing to the degradation of the lake.

Determination of Hydraulic Head

The total shoreline of Austin Lake was divided into segments. Most lots represented one segment. Large lots were divided into two or more segments. Areas not designated as lots were divided into segments that were approximately 50 feet in width. Each of the 678 segments was assigned a number. Fifty segments were selected at random for the study (Figure 13).

In order to determine the direction of ground-water flow, minipiezometers were placed at the approximate center point of each segment at distances of 10 feet (3.0 m) and 20 feet (6.1 m) from shore.

The miniplezometer, modified from one used by Lee and Cherry (1978), consists of a 3/4 inch (1.9 cm) iron

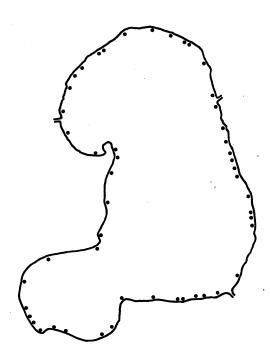


Figure 13. Approximate locations of sample sites on Austin Lake.

pipe. A 1/2 inch (1.3 cm) iron pipe 8 inches (20.3 cm)long was welded to one end. A 7/16 inch (1.1 cm) polyethylene tube 8 inches (20.3 cm) long was placed inside the 1/2 inch (1.3 cm) iron pipe and the two were drilled simitaneously and perforated with a 1/4 inch (0.6 cm) bit. The polyethylene tube was then removed, wrapped with a fine mesh nylon net and replaced. Care was taken to be sure the holes on the pipe and the tube lined up. They were then fixed in place with a small bolt through the pipe and tube. The 1/2 inch (1.3 cm) pipe was capped with a reducer and plugged. The reducer and plug served as a driving point.

The pipe was capped and driven to the desired depth. The cap was then removed and water allowed to rise up into the pipe. To facilitate the process, water was added until the water level in the pipe was at the level of the water in the lake. Subsequent water-level movement within the pipe was measured and recorded along with water depth and pipe depth into the sediment. Waterlevel movement was checked by varying the level of water in the pipe to determine if the results could be replicated. The difference between the water levels of the lake and pipe is the hydraulic head. To facilitate the measurement of the hydraulic head, a device (manometer) was inserted into the minipiezometer. Readings could then be made above the surface of the lake (Figure 14).

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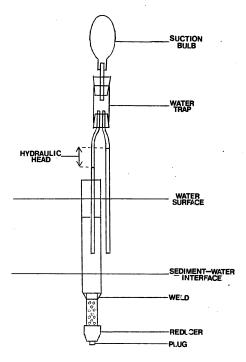


Figure 14. Sampling equipment.

Water Sampling

The miniplezometer was pumped dry (not always possible in permeable sediments) and ground-water samples taken from water that refilled the pipe from below.

Chemical Analysis

Each sample was frozen at the end of each sampling day. When enough samples accumulated, they were thawed. The specific conductivity was determined and the sample analyzed for its chloride and phosphate content.

Conductivity was measured using model EP-10 specific conductivity meter manufactured by the Myron L. Company of Encinitas, California. Results are given in micromhos/centimeter.

Chloride concentration was determined using the Hach mercuric nitrate method. In this procedure, diphenylcarbazone is the indicator and mercuric nitrate the titrant. Once all the chloride present in the sample is complexed by the mercuric ions, the excess mercuric ions combine with diphenylcarbazone to form a purple complex indicating the end point of the titration.

Phosphate concentration was determined using the Hach ascorbic acid method. In this phosphate analysis, orthophosphate reacts with molybdate in an acidic medium to produce phosphomolybdic acid that is then reduced to

a heteropoly blue compound. The phosphate concentration of the water sample is then determined by measuring the intensity of the blue color.

Importance of Water Quality Parameters Measured

Conductivity is a measure of the ability of water to conduct an electrical current. It is expressed in micromhos per centimeter at 25°C. Pure water has a low conductivity. Conductivity increases with increasing concentrations of dissolved minerals in the water. Normally the amount of dissolved solids in milligrams per liter (mg/l) is approximately 65 percent of the conductivity. Dissolved solids are often contributed to ground water by industrial and domestic wastes.

Soluble Orthophosphate--Phosphorus is the dissolved form of phosphate that is directly available for plant growth. Phosphate is reported rather than phosphorus because the element, phosphorus, does not occur free in nature. Instead, it is found in combined forms, the most common being phosphates. Phosphorus, like nitrogen, is an element essential to plant growth and both elements can be derived from sources such as soils and decomposition of organic matter. Phosphorus has become increasingly evident in the environment because of its past extensive use in detergents and its present use in agriculture. In streams, phosphorus is an extremely limiting

element for plant growth and thus an increased supply from outside sources such as sewage effluents may provide the stimulus for excessive production.

Phosphate movement in ground water is limited and therefore was not expected to be detected in large concentrations. Comparisons of phosphate concentrations with chloride concentrations may yield additional information regarding these species.

The chloride ion, in combination with one of several cations such as sodium, magnesium, and potassium, is a common impurity in most natural waters. Chloride compounds are readily dissolved from rocks by surface and subsurface waters. A concentration of several hundred milligrams per liter is necessary before a salty taste is imparted to the water or aquatic organisms are harmed.

Chlorides are essential in the diet and passes through the digestive system to become one of the major components of sewage. The wide use of salt in water softeners also contributes a large amount of chloride to sewage. In water, chloride concentrations are attenuated primarily by dilution and therefore high concentrations are good indicators of pollution.

CHAPTER 4

ANALYSIS OF DATA

The data were analyzed on an IBM 3033 computer operating from an IBM 3270 terminal under Time Sharing Option (TSO). The language used throughout the analysis was Statistical Analysis System (SAS).

Table Descriptions

The data analyzed are listed in Tables 1, 2, and 3.

Table 1, on pages 36, 37, 38, and 39, contains physical data from each sampling point. The data in Table 1 include the depth of the water at the sampling point (DW), the depth the piezometer was driven into the sediment (DP), the hydraulic head (HH), and whether or not a functioning septic tank was present on the lot onshore from the sampling site.

Table 2, on pages 40, 41, 42, and 43, contains the results of the ground-water sample analyses for each sampling point. The data in Table 2 include specific conductivity (COND), chloride content (CHL), and phosphate content (P).

Table 3, on pages 44, 45, and 46, contains data on other physical factors that may have an effect on groundwater movement and/or chemical properties of the ground

Location Number		of Water m)	Depth o in Sedim	f Pipe ent (cm)	Hydraul: (m		Septic Tank
Number.	3 m	6 m	3 m	6 m	3 m	6 m	(yes, no)
7	-	35	-	56	_	-5	no
18	-	-	61	61	-35	-10	yes
, 32	25	61	61	46	0	-10	yes
49	-	20	-	61	-	-65	yes
68	76	-	61	-	+7	-	yes
75	17	35	57	47	+3	+5	yes
120	34	96	55	59	- - 155	-189	yes
144	79	152	87	61	+2	-9	yes
168	62	122	65	55	-190	-280	yes
170	28	51	51	87	0	0	no
176	23	44	47	121	0	o	no
185	51	98	52	123	-3	-5	no

Table 1	
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Physical Data for Each Study Site

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Table 1 (Continued)

yes. no) Septic ank ou res 0 H g yes оц 8 80 yes g res yes res -215 -100 -120 -110 -125 <u></u>2 -45 Hydraulic Head Ħ 29 29 42 4 1 ŝ 6 100 5 8 ĥ E 29 29 22 0 ĥ 5 m (m) Ξ Depth of Fipe n Sediment (cr 5 8 95 95 95 105 104 \$ 9 g 92 9 102 ø Ħ 113 ß 89 5 20 95 2 65 M ß Depth of Water (cm) 6m 38 ß 2 ß 20 38 5 20 48 ŝ 5 \$ E 4 28 g 20 20 35 6 ß 88 m ы g 33 5 Location Number 305 238 244 260 274 281 287 320 192 23 275 361 365

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Table 1 (Continued)

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Septic Tank	(yes, no)	yes	yes	ou	ou	ou	ou	ou	ou	ou	ou	ou	yes	Ves
ic Head m)	н 9	-40	-25	+5	0	0	-5	-15	0	-110	0	6-	-180	-200
Hydraulic (mm)	а 6	-42	-10	0	0	0	0	-25	0	-75	0	5	-150	-170
Depth of Pipe in Sediment (cm)	е 9	67	98	95	92	95	101	109	56	88	64	55	86	5
Depth in Sedi	ы Н	20	67	. 98	68	68	81	53	50	86	76	45	Z	68
Depth of Water (cm)	ы Ф	28	52	38	38	. 33	2	43	46	64	104	67	2	.15
	E E	12	28	30	30	30	Z	38	38	41	٣	43	38	38
Location	Number	366	373	383	401	403	416	419	420	451	457	473	496	497

Table 1 (Continued)

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	and the second se		And a second sec				
Location	Depth of Water (cm)	Water 1)	Depth of Pipe in Sediment (c	f Pipe ent (cm)	Hydraulic (mm)	Head	Septic Tank
Number.	Э ш	Ш 9	ы Ш	ш 9	Э Ш	н 9	(Yes, no)
508	41	74	42	8	-255	-240	yes
519	2	74	119	73	-3	۲	ou
548	58	66	74	53	-5	-10	ou .
560	41	69	88	83	-230	-360	ou
563	ı	ı	75	8	ı	-220	ou
579	I	61	83	16	ı	-310	ou
585	41	I	63	96	۰ ۲	ı	ou
596	35	53	35	116	-1-	1	ou
614	66	74	56	48	42	0	ou
650	54	39	76	66	+70	+10	no
663	45	2	59	96	+28	-30	ou

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Table 2

Water Analysis Results for Each Study Site

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		AT AT A PTOTT TOOD	101 0010000	00 T 2 / 200 200 T 10 T	2	
Location Number	Specific Conductivit (Wmhos/cm)	fic ivity v/cm)	Chlorides (mg/l)	des (1)	Phosphates (mg/l)	lates /1)
	3 m	е п	З ш	е ш	3 ш	ш 9
2	ł	320	1	30	I	0.08
18	335	340	30	30	0.03	0.03
32	320	350	32	32	60.0	0.1
6†	ı	380	I	40	ı	60.0
68	620	ı	20	ı	0.1	ı
75	305	325	35	35	60°0	0.06
120	340	360	35	32	1.0	0.55
144	320	320	27	30	0.08	0.1
168	320	360	35	30	0.5	0.5
170	280	ı	30	1	0.15	ı
176	320	. 420	30	- 32	0.25	0.25
185	310	044	30	30	1.4	1.15

40

Table 2 (Continued)

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Location Number	Specific Conductivi (/mhos/cm	fic ivity /cm)	Chlorides (mg/l)	des (1)	Phosphates (mg/1)	ates 1)
	5 ш	6 ш	3 ш	6 m	3 m	е 9
192	270	270	30	30	0.15	0
213	360	280	35	30	ı	2.5
218	ı	280	35	. 40	ı	0.2
544	320	360	35	30	0.6	0.25
260	320	310	30	30	0.4	0.7
274	ı	ı	ı	ı	ł	ı
275	380	400	30	30	6.0	0.35
281	320	280	30	30	0.25	0.25
287	310	340	30	30	0.25	0.3
305	260	260	30	30	0.15	0.2
320	280	320	30	30	0.15	0.15
361	044	280	30	30	0.2	0.25
365	360		25	45	0.25	0.25

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Table 2 (Continued)

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Ð	Specific Conductivi (\umbos/cm	ific tivity s/cm)	Chlor (mg	Chlorides (mg/l)	Phosp (mg	Phosphates (mg/l)
• • • •	а п	9 9	3 m	6 m	5 m	е 9
	320	280	30	35	0.4	0.2
ন	400	260	0†	30	0.2	0.25
N)	320	260	30	30	0.15	0.15
ГМ	360	320	30	30	0.15	0.13
~	360	360	30	25	0.15	0.4
~ `	320	320	30	30	0-05	0.3
гл	320	310	35	30	0.05	0.1
м	320	290	25	25	0.1	0.1
М	300	300	30	25	0.1	0.3
R)	330	300	25	30	0.1	0.1
N)	300	320	30	25	0.1	0.1
~	300	340	25	25	0.1	0.1
NO.	320	360	25	30	0	0.15

42

Table 2 (Continued)

Location Number	Spec Conduc (Pmho	Specific Conductivity (Mmhos/cm)	Chlorides (mg/l)	des 1)	Phosphates (mg/l)	ates 1)
	З ш	е ш	5 m	6 m	Э ш С	ш 9
508	300	300	30	30	0.1	0.1
519	320	360	35	30	0.25	0.25
548	360	400	30	30	0.8	0.5
560	270	270	30	30	0.1	0.15
563	280	260	25	30	1	0.3
579	260	250	25	30	0.1	0.15
585	300	ı	25	ı	0.05	ı
596	300	280	25	30	0.1	0.2
614	340	330	25	30	0.2	0.3
650	380	360	30	35	0.25	0.2
663	410	540	40	40	0.1	0.15

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43

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Location	Drainage Basin	Soil ^a	Glacial Material ^b	Permea	bilityc
Number	Altitude (ft)		Material ^b	3 m	6 m -
7	859	02	0	-	3
18	861	02	0.	2	2
32	864	02	0	1	1
49	868	03	0	2	2.
68	868	03	0	2	-
75	868	03	0	1	1
120	876	01	0	3	1
144	868	01	0	2	3
168	872	03	0	3	3
170	870	03	0	2	2
176	860	03	0	2	3
185	861	03	0	2	2.
192	870	03	0	2	2
213	868	04	0	3	3
218	868	03	0	3	3
244	869	04	0	3	3
260	868	04	0	3	3
274	864	04	0	3	3
275	864	04	0	2	3
281	863	04	0	3	3

Table 3 Factors Considered in the Statistical Analysis of Austin Eake's Ground-water Characteristics

Location Number	Drainage Basin Altitude	Soil ^a	Glacial ^b Material		bilityc
Number	(ft)			3 m	6 m
287	862	04	0	3	3
305	858	A	С	2	2
320	861	в	С	2	2
361	861	04	C	3	3
365	861	04	С	3	3.
366	861	04	С	3	3
373	861	04	C	3	3
383	861	04	С	2	2
401	861	04	С	3	3
403	861	04	С	3	2
4 1 6	861	04	C	3	3
419	861	04	C	3	3
420	861	04	C	3	3
451	859	04	C	3	3
457	859	04	C	3	3
473	859	04	М	3	3
496	861	04	М	2	2
497	861	04	М	2	2
508	86 1	04	М	2	2
519	861	04	Μ	3	2
548	858	04	0	2	2

Table 3 (Continued)

Location Number	Drainage Basin Altitude (ft)	Soil ^a	Glacial _b Material	Permea 3 m	bility ^C 6 m
560	861	04	0	1	1
563	862	04	0	-	2
579	865	04	0	1	1
585	870	04	0	3	3
596	868	04	0	3	3
614	859	A	0	3	3
650	868	02	0	3	3
663	861	A	0	3	2

Table 3 (Continued)

aO1, 02, 03, 04 = varieties of Ostemo
A = Adrian
B = Brady

 $b_0 = outwash$

M = morainal material C = channel deposits

 $c_1 = excellent$ 2 = good 3 = poor

water. These factors are the altitude of the drainage basin (DBA), the type of soil adjacent to and surrounding the lake, the surface glacial material adjacent to and surrounding the lake, and a subjective judgment of the permeability of the sediment at each sampling point.

The altitude of the drainage basin is the maximal altitude in the drainage basin perpendicular to the shoreline at the sampling site. The types of soil adjacent to and surrounding the lake include four varieties of Oshtemo (01, 02, 03, 04), Brady (B), and Adrian (A). The types of surface material adjacent to and surrounding the lake are channel deposits (C), outwash (O), and morainal material (M). Permeability is classified as excellent (1), good (2), and poor (3).

Variable Descriptions

Variables are classified according to whether they are measurement or classification variables and whether they are dependent or independent variables.

Measurement variables have measured values. Classification variables do not have measured values, or if they do, they can be simply grouped and given a group name. Hydraulic head, depth of the water, depth of the piezometer in the sediment, altitude of the drainage basin, specific conductivity, chloride content, and phosphate content are all measurement variables. The classification

variables are glacial material, soil, permeability of sediment, distance from shore, and presence or absence of a septic tank.

Dependent variables are subject to change if conditions change. Independent variables are fixed. They do not change. Hydraulic head, specific conductivity, chloride content, and phosphate content are dependent variables. Independent variables include all the classification variables.

Analytical Procedure

The first step in the analyses was to examine interdependencies among the variables DP, COND, HH, CHL, and P using pairwise Pearson Product Moment correlations (PROC CORR). In this procedure correlations were calculated considering breakdowns involving the independent variables--septic tank and distance from shore in the following manner: (1) all observations combined (Appendix A); (2) observations classified according to values on both independent variables (Appendix B); and (3) observations classified according to the presence or absence of a septic tank only (Appendix C).

The second step in the analyses used two-way analysis of variance procedures to determine the relationship between septic tank and distance from shore on the dependent variables HH, COND, CHL, and P. Because of unequal

sample sizes a general linear models program (GLM) was used for this analysis. This analysis was performed in the following manner: (1) observations classified according to values on both independent variables (Appendix D) and (2) observations classified according to the presence or absence of a septic tank only (Appendix E).

Schematic plots were drawn for each dependent variable to determine if visual inspection could identify any differences in the effects of a septic tank and/or distance from shore (Appendix F). Examination of the plots indicated that other factors might be related to the results. Factors considered were the altitude of the drainage basin, depth of water at each sampling point, the soil surrounding the lake, and the surface glacial material surrounding the lake. Those parameters were added to the data base along with a subjective determination of the permeability of the sediment at each sampling site (Table 3).

The dependent variable data were examined for skewness using the univariate procedure. Log transformations were made on the data to reduce the skewness for all dependent variables except hydraulic head (Appendix G).

In the final analyses the GLM procedure was used because both classification and measurement variables were among the independent variables. These analyses included the added factors and the log transformations.

49

In the final analysis of HH, the classification variables were types of glacial material, types of soil, and permeability of the sediment. The measurement variables were the altitude of the drainage basin, depth of the piezometer, and the depth of the water (Appendix H).

In the final analysis of the other dependent variables (COND, CHL, P) the classification variables were septic tank, distance from shore, glacial material, soil, and permeability. The measurement variables were hydraulic head, altitude of the drainage basin, depth of the piezometer, and depth of the water (Appendix I).

In all analyses using the GLM procedure, the results referred to are from using the Type IV sum of squares. The Type IV sum of squares is the most appropriate for this study because it is adjusted for differences in sample size.

Computer Results

Hydraulic Head

In the initial correlational procedure observations were not grouped according to the independent variables septic tank and distance from shore. Here hydraulic head was seen to have a significant relationship to conductivity, \underline{r} (84) = 0.21843, $p \le 0.0433$. At locations without a septic tank, there was a highly significant relationship

between hydraulic head and conductivity at a distance of 3 meters, \underline{r} (25) = 0.49835, $p \le 0.0082$. At 6 meters, the relationship of hydraulic head to conductivity was marginally significant, \underline{r} (24) = 0.36871, $p \le 0.0638$. At locations with a septic tank the relationship of hydraulic head to conductivity was marginally significant at 3 meters, \underline{r} (14) = 0.43058, $p \le 0.0959$. At 6 meters from shore, the relationship of hydraulic head to conductivity was not found to be significant, but the relationship of hydraulic head to phosphates was significant, \underline{r} (15) = -0.48480, $p \le 0.0486$. Because the relationship of hydraulic head to phosphates was shown to be significant in only one correlation set, it is likely that this statistic is the result of chance alone.

In the second series of correlations in which distance from shore was not a factor, the relationship of hydraulic head to conductivity was significant to localities where a septic tank was not present, $\underline{\mathbf{r}}$ (51) = 0.34441, $\mathbf{p} \leq 0.0116$. When a septic tank was present, the relationship of hydraulic head to phosphates was marginally significant, \mathbf{r} (31) = -0.32440, $\mathbf{p} \leq 0.0655$.

The GLM procedure showed an overall significant relationship to hydraulic head, $\underline{F}(3, 86) = 3.54$, $p \le 0.0179$. Individually there was a significant relationship of hydraulic head to the presence of a septic tank, $\underline{F}(1, 86)$

= 5.14, $p \le 0.0259$, and a significant relationship of hydraulic head to distance from shore, <u>F</u> (1, 86) = 4.31, $p \le 0.0408$. The interaction between septic tank and distance from shore for hydraulic head was not found to be significant, <u>F</u> (1, 86) = 0.41, $p \le 0.5228$.

The schematic plot showed that the hydraulic head decreased with distance from shore and decreased in the presence of a septic tank.

In the final analysis of hydraulic head, the independent variables were the class variables (glacial material, soil, and permeability) and the measurement variables (altitude of the drainage basin, water depth, and the depth of the piezometer). This model was highly significant, \underline{F} (12, 74) = 4.0, p \leq 0.0001, and explained about 39 percent of the variability.

Overall, the relationship of hydraulic head to the type of glacial material was significant, \underline{F} (2, 74) = 3.87, $p \neq 0.0251$. The least squares means for channel deposits and outwash were -52.8 and 49.5 respectively and the least squares mean for morainal material was -130.6. This indicates a greater outflow in the area of the lake that is bounded by morainal material.

The relationship of hydraulic head to soil type was marginally significant, \underline{F} (5, 74) = 2.10, p \leq 0.0738. Because of this overall marginal significance little emphasis is placed on it.

The relationship of hydraulic head to the permeability of the sediment was highly significant, <u>F</u> (2, 74) = 9.62, $p \le 0.0002$. From the least squares means, poor permeability, with a least squares mean of -30.19, was associated with the highest hydraulic head (most positive number and lowest outward flow) as compared to excellent permeability with a least squares mean of -159.0.

The relationship of hydraulic head to the altitude of the drainage basin was significant, $\underline{F}(1, 74) = 5.33$, $p \le 0.0237$. The estimate of the linear coefficient on the effect of the drainage basin altitude is -22.66. This is the slope of the linear regression line and being negative gives a result of a more negative head with an increase in the altitude of the basin.

The relationship of hydraulic head to the depth of the piezometer in the sediment was significant, \underline{F} (1, 74) = 4.32, p \leq 0.0411. The estimate of the linear coefficient on the depth is -0.80. This means that the deeper the piezometer is placed in the sediment, the greater the negative hydraulic potential differential.

The relationship of hydraulic head to water depth was significant, $\underline{F}(1, 74) = 4.29$, $p \le 0.0419$. The estimate of the linear coefficient on the water depth is -0.77. This is negative, therefore the greater the depth, the greater the negative hydraulic potential differential.

Conductivity

In the initial correlational procedure that did not group observations according to the independent variables septic tank and distance from shore, the correlation of conductivity with hydraulic head was significant as reported in the previous section. In addition, when no septic tank was present, conductivity had a significant relationship to chlorides at 3 meters, \underline{r} (27) = 0.38365, p \leq 0.0399, and a highly significant relationship to chlorides at 6 meters, \underline{r} (26) = 0.53692, p \leq 0.0032.

In the second series of correlations when distance from shore was not considered in the analysis, again conductivity had a significant relationship to hydraulic head as reported previously. Conductivity had a very high significant relationship to chlorides, \underline{r} (55) = 0.43355, $p \neq 0.0008$, when no septic tank was present. With the presence of a septic tank, conductivity had a marginally significant relationship to the depth of the piezometer, r (31) = -0.30350, $p \neq 0.0860$.

The schematic plots for conductivity showed a marginal increase in the presence of a septic tank. There was a marginal increase with distance when the septic tank was absent and a decrease with the presence of a septic tank.

In the final analysis, the independent class variables

are glacial material, soil type, permeability, distance from shore, and septic tank. The independent measurement variables are hydraulic head, altitude of the drainage basin, depth of the piezometer in the sediment, and depth of the water.

The full model for the log of conductivity was not found to be significant, \underline{F} (15, 67) = 1.17, p 4 0.3151. In this model conductivity had a marginally significant relationship to hydraulic head, \underline{F} (1, 67) = 3.19, p 4 0.0788. The estimate of the linear coefficient for the effect of the hydraulic head was 0.0004. This is positive and means that the greater the hydraulic head the greater the conductivity. In other works, the faster the water is flowing out of the lake, the lower the conductivity.

None of the other variables was found to be significant and these variables are not considered further.

A reduced model was designed for the log of conductivity. This model contained only the factors that showed significant values in the full model. These factors were hydraulic head and septic tank. This model was significant, \underline{F} (3, 82) = 3.29, p \leq 0.0246, and explained almost 11 percent of the variability. In this model conductivity had a significant relationship to hydraulic head, \underline{F} (1, 82) = 6.01, p \leq 0.0164. The presence of a septic tank no longer showed a particularly strong relationship, but when considered separately from hydraulic head, the

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relationship became marginally significant. This implies that the conductivity is greater in the presence of a septic tank.

Chlorides

In the correlational procedures, significant relationships were not found other than that of conductivity previously mentioned.

The GLM procedure showed a marginal significant relationship between chloride content and septic tank presence, F (1, 87) = 3.22, p ≤ 0.0761 .

The schematic plot for chlorides showed a fairly strong relationship with the presence of septic tanks. Chloride content was greater and varied more in the presence of a septic tank.

In the final analysis, the independent variables were the same as in the analysis of the log of conductivity. The full model for the log of chlorides was not found to be significant, \underline{F} (15, 68) = 1.53, p 4 0.1198. Chlorides had a marginally significant relationship to permeability, \underline{F} (2, 68) = 2.38, p = 0.1001, and a significant relationship to the depth of the water, \underline{F} (1, 68) = 4.13, p 4 0.0459. The estimate of the linear coefficient for water depth was negative (-0.0016), therefore the deeper the water, the lower the level of chlorides. None of the other relationships was strong enough to consider.

Phosphates

The initial correlations did not reveal any significant relationships between phosphate content and the several variables.

In the full model for the log of phosphates, the independent variables were the same as in the analyses of the other full models. The overall model was not found to be significant, \underline{F} (15, 66) = 1.27, p \pm 0.2450. Glacial material had the only significant relationship to phosphate content, \underline{F} (2, 66) = 3.75, p \pm 0.0288. The least squares mean for outwash is 0.242 as compared to channel deposits at 0.074 and morainal deposits at 0.004. This means that outwash has a greater level of phosphates, but overall, phosphate content was not found to be significant.

CHAPTER 5

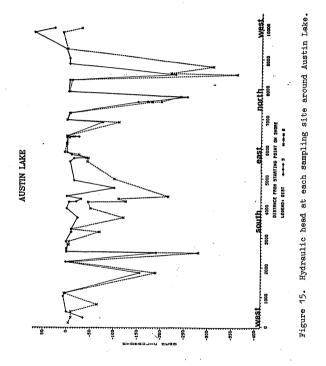
DISCUSSION AND CONCLUSIONS

Ground-water Flow

The first objective of this investigation was to determine the direction of ground-water flow through the sediment-water interface along the total shoreline of Austin Lake. This was done by placing minipiezometers in the sediment beneath the water surface and measuring the hydraulic head. The data obtained show that at most sites the water movement has a downward component. The greatest downward flow appears to be at the northern, southern, and southwestern portions of the lake (Figure 15). Few measurements show ground water moving into the lake. The compass direction of flow cannot be determined from the data, but can be inferred from previous data as presented by Deutsch et al. (1960) and Allen et al. (1972). Their models suggest that ground water moves to the north and to the south from Austin Lake. Austin Lake is situated on the ground-water divide.

During the collection of hydraulic head data, three conditions were found to exist: (1) a negative hydraulic head for both samples at the sampling site; (2) a positive hydraulic head for both samples at the sampling site; and

58



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(3) a zero or a positive hydraulic head at 3 meters and a negative hydraulic head at 6 meters.

In the first situation, water is seeping through the lake seal and becoming part of the ground-water flow (recharge point). In the second situation, the ground-water movement is upward and into the lake (discharge point). The third situation suggests a combination of the first two. Ground water is moving into the lake at 3 meters and moving out of the lake at 6 meters (Figure 16). This last situation could occur if there is a slight local elevation of the water table near shore and would be enhanced by the presence of a channel for ground-water movement into the lake. These channels may be present. During sampling it was noticed that the sediment was layered at many points around the lake. These layers were narrow--from one to several centimeters in thickness. They were made of sand interbedded with finer material with organic material scattered throughout. At some of these sites the piezometer was driven deeper. The hydraulic head then became negative.

Computer analysis shows several factors affecting the hydraulic head. One factor is the depth of the piezometer in the sediment. This was expected. The deeper the piezometer is in the sediment, the more equipotential lines intersected, the greater the hydraulic head (Figure 8).

Another factor affecting the hydraulic head is the

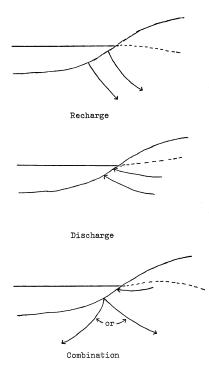


Figure 16. Possible ground-water flow patterns at Austin Lake.

depth of the water. As the depth of the water increases, the hydraulic head becomes more negative. This statistic is probably the result of the greatest outward flow being at the areas of the lake that were the deepest, namely the northern, southern, and southwestern portions.

The permeability of the sediment affects the hydraulic head. The largest hydraulic head differential occurs where the permeability is the greatest. At one sampling site that exhibited excellent permeability, a siphoning hose was placed from the lake into the minipiezometer. Water ran into the piezometer (out of the lake) for some time until the hose was removed. This indicates a thin seal in the lake. Once through the seal, the sediments are permeable.

The statistical analysis of the relationship of drainage basin altitude to hydraulic head is significant but not in the way expected. The higher the altitude of the drainage basin, the more negative the hydraulic head. Examination of the sampling results show high negative hydraulic head (outward) readings in the southwestern region of the lake. This region contains some of the highest points in the drainage basin. The southwestern region is illustrated on the cross-section of Austin Lake (Figure 2). The left side of A-A' is the surface of Gourdneck Lake. This surface of Gourdneck Lake is below the surface of Austin Lake. It is reasonable to assume

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that this lower surface has a major effect on ground-water flow from Austin Lake towards Gourdneck Lake and that the topographic high between the two lakes is overshadowed by this effect.

The northern shore of Austin Lake also shows a large negative head. The surface-water divide is to the north (Figure 6). This neans that surface water flows to the south and ground water flows to the north in this region. Therefore altitude of the drainage basin has a minimal effect on the total ground-water flow at Austin Lake.

In glacial sediment, morainal material shows the greatest outward flow (Figure 17). It is unclear whether this is a true picture of what is happening because of the small extent of morainal material adjoining the lake. These glacial materials are surface deposits and may have only a minimal effect on regional ground-water flow.

Pollution

The second objective was to determine if the lake is being polluted by ground water by examining the specific conductivity, chloride content, and phosphate content of ground-water samples from beneath the lake.

Three factors, hydraulic head, depth of the piezometer in the sediment, and presence or absence of a septic tank, affect the specific conductivity. The more positive the hydraulic head, the greater the conductivity (Figure 18).

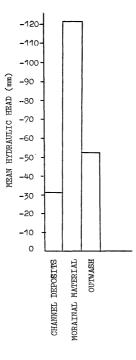
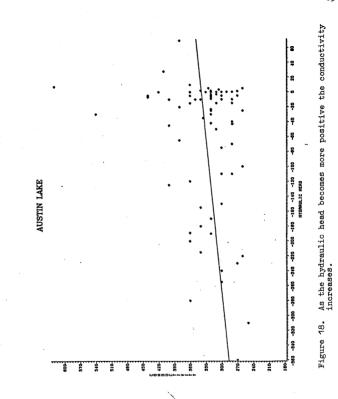


Figure 17. Average hydraulic head readings for the different types of glacial material.

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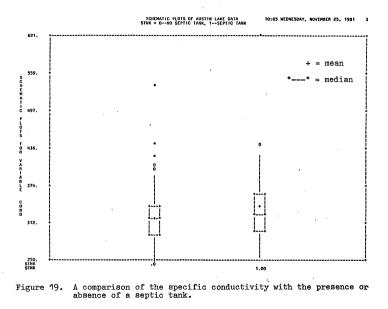
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In other words, the faster water is moving out of the lake, the lower the conductivity. The deeper the piezometer was placed in the sediment, the greater the conductivity. Conductivity is higher when a septic tank is present (Figure 19). Consequently, it appears that septic tank effluent is entering the lake. These results seem to indicate that at certain times during the year. at some points and under certain conditions, ground water moves into the lake bringing in dissolved compounds. Later when the ground-water direction is reversed, the compounds are carried downward and away from the lake, therefore the higher concentrations will be at the lower level. This idea is strengthened by the fact that a late summer check of the hydraulic head on the west side of Austin Lake indicated an outward flow even though some of the points had indicated an inward flow earlier in the spring and summer.

Chlorides exhibit some of the same relationships as conductivity. Offshore from septic tanks chloride content varies considerably. At locations without a septic tank chloride content varies little (Figure 20). This difference in variance indicates some inflow of effluent from septic tanks.

Significant relationships were not found between phosphates and any of the dependent variables. This is

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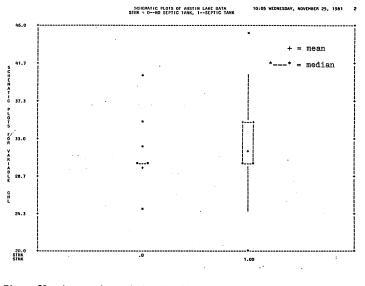


Figure 20. A comparison of the chloride concentrations with the presence or absence of a septic tank.

as expected as phosphates are rapidly attenuated in sediment. One interesting result is that a greater concentration of phosphates was present near glacial outwash. This area contains mostly older homes and the possibility exists that phosphates are entering the lake from older functional or non-functional septic systems (Figure 21).

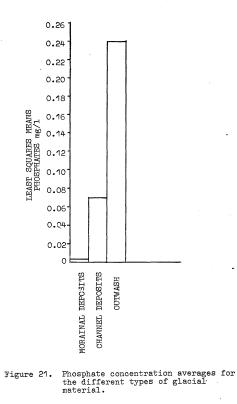
Summary of Conclusions

The conclusions directly related to the objectives are:

- The primary direction of water flow is downward through the sediment-water interface.
- Under certain conditions ground water moves up through the sediment-water interface.
- Upward movement of ground water brings in dissolved materials from onshore septic systems.
- The amount of septic tank effluent entering the lake is minimal.
- The lake-bottom sediment forms a thin seal that retards, but does not prevent, downward movement of water.

Conclusions not related to the objectives but that are consistent with those of other researchers are:

 The deeper the piezometer is in the sediment, the greater the hydraulic differential.



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 Water movement is primarily north and south of Austin Lake, with little east or west flow. This suggests that the divide is along an eastwest line that crosses Austin Lake and probably changes position during the year as the lake varies from recharging to discharging at particular points.

CHAPTER 6

RECOMMENDATIONS

Recommendations for future investigations include:
A study of the near shore hydrology. What are the effects of the layered sediments on local ground-water flow? Could these layers shed any light on historical development of the lake?

- 2. A study of conditions necessary for water movement into the lake. Does water move into the lake at all points or is it limited in extent-near topographic highs or near the ground-water divide on the east and west sides?
- 3. A study of the water movement through the organic sediment filled basins. Are the basins impermeable? Do they act as sinks for dissolved materials?

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APPENDIX A

PEARSON PRODUCT MOMENT CORRELATIONS OF ALL OBSERVATIONS COMBINED

	OBSERVATIONS	CORRELATION COEFFICIENTS / PROB > IR UNDER HO:RHO=O / NUMBER OF OBSERVATIONS	IENTS / PROB > IRI UN	CORRELATION COEFFIC		
2.5000000	•	22.9800000	0.33696561	0.26113636	88	•
45.000000	20.0000000	2770.0000000	4.05846574	30.43956044	16	CHL
620,0000000	250.0000000	29585.0000000	57.30149473	328.7222222	8	COND
70.000000	-360.0000000	-4733.0000000	86.02702664	-52.5888889	90	1
123.000000	35.0000000	7290.0000000	20,96743625	76.73681211	56	40
UNIXYN	NUMININ	NNS	STD DEV	нели	-	VARIABLE
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CHL 0.10645 0.3152 0.6112 0.6112 0.12291 0.12291 0.2229 0.23506 0.05506 0.21642 0.8055 0.0435 0.0435 0.0435 0.0000 0.12991 0.12991 0.12991 0.12991 0.12291 0.02352 0.02352

-0.12954 -0.2236 -0.00000 0.0000 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0133 0.0135 0.0155 0.0135 0.0135 0.0135 0.0135 0.0135 0.0015 0.0135 0.0015 0.

0.100000 0.22554 0.22554 0.22554 0.22554 0.10554 0.10554 0.10554

APPENDIX B

FEARSON PRODUCT MOMENT CORRELATIONS OF OBSERVATIONS CLASSIFIED ACCORDING TO VALUES ON INDEFENDENT VARIABLES-DISTANCE FROM SHORE AND PRESENCE OR ABSENCE OF A SEPTIC TANK

	OBSERVATIONS	CORRELATION COEFFICIENTS / PROB > [R] UNDER HO:RMO=0 / NUMBER OF OBSERVATIONS	cients / Prob > Irl un	CORRELATION COEFFIC		
1.4000000	0.0500000	6.5000000	0.28623452	0.24074074	27	•
00000000.04	25.0000000	865.0000000	3.65541639	29.82758621	29	CHL
410.000000	260.0000000	9180.0000000	33.94069008	316.55172414	53	COND
70.0000000	-230.0000000	-333.0000000	49.06980899	- 12. 33333333	27	H
119.000000	35.0000000	2005-0000000	20.61331771	69.13793103	53	40
NUNIXYN	MUMININ	HNS	STD DEV	нели	Ŧ	VARIABLE
		DIST=10	STNK=0			

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-0.06973	0.09592	0.11225	0.13145	1.00000
0.7297	0.6411	0.5772	0.5134	0.0000
27	26	27	27	27
0.17333	0.01702	0.38365	1.00000	0.13145
0.3686	0.9328	0.0399	0.0000	0.5134
29	27	29	29	27
-0.00287	0.49835	1.00000	0.38365	0.11225
0.9882	0.0082	0.0000	0.0399	0.5772
29	27	29	29	27
-0.15327	1.00000	0.49835	0.01702	0.09592
0.4453	0.0000	0.0082	0.9328	0.6411
27	27	27	27	26
1.00000	-0.15327	-0.00287	0.17333	-0.06973
0.0000	0.4453	0.9882	0.3686	0.7297
29	27	29	29	27

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			SINK=0	STATESTEALANALYSIS SYSTEM SIMK=0 DIST=20	H 18:20 FRIDAY, OCTOBER 23, 1981	JOBER 23, 1981
VARIABLE	z	HEAN	STD DEV	NNS	MUMININ	INHIXVH
5	30	86.5666667	20.25643079	2597,0000000	46.000000	123.0000000
Ħ	27	-60.62962963	101.61006938	-1637.0000000	-360.0000000	10.000000
COND	28	327.5000000	63.34064285	9170.0000000	250.0000000	540.000000
CHL	28	29.89285714	2.91025372	837.0000000	25.0000000	40.000000
•	28	0.33607143	0.48096615	9.4100000	•	2.500000
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4	1.00000 0.0000 30	-0, 12863 0.5226 27	0.22496 0.2498 28	0.27730 0.1531 28	0.1948
Ŧ	-0.12363 0.5226 27	1.00000 0.0000 27	0.36871 0.0638 26	0.01878 0.9275 26	0.0188
COND	0.22496 0.2498 28	0.36871 0.0638 26	1.00000 0.0000 28	0.53696 0.0032 28	0.929
CK	0.27730 0.1531 28	0.01878 0.9275 26	0.53696 0.0032 28	1.00000 0.0000 28	-0.0394 0.8419 21
•	0.19485 0.3204	0.01681 0.9273 26	0.01754 0.9294 28	-0.03947 0.6419 28	0.0000

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VARIABLE	Ŧ	нели	STD DEV	NNS	NUMININ	NUHIXAN
90	81	70.1666667	16.21999492	1263.0000000	51.0000000	113.0000000
Ŧ	2	-61,.0000000	82.87907563	-1152.0000000	-255.0000000	1.0000000
COND	16	352.5000000	83.40663443	5640.0000000	260.0000000	620.00001000
CHL	11	30.23529412	4.91845265	514.0000000	20.0000000	40.0000000
•	91	0.22437500	0.24671086	3.5900000	0	1.0000000
		CORRELATION COEFFIC	IENTS / PROB > IR! UNC	CORRELATION COEFFICIENTS / PROB > R UNDER HO:RHO=0 / NUMBER OF OBSERVATIONS	SERVATIONS	
			DP HH	COND CHL P		
N.		90	1.00000 -0.07207 -	1.00000 -0.07207 -0.29363 -0.18819 -0.24672 0.0000 0.7763 0.2697 0.4695 0.3570		

•	-0.24672 0.3570	-0.25374 0.3430 16	-0.04074 0.6809 16	0.40456 0.1201 16	0.0000
R	-0.18819 0.4695 17	0.00134 0.9959 17	-0.36439 0.1653	0.0000	0.40456 0.1201 16
COND	-0.29363 0.2697	0.43056 0.0959	0.0000	-0.36439 0.1653 16	-0.04074 0.8809 16
Ŧ	-0.07207 0.7763	1.00000 0.0000 18	0.43058 0.0959 16	0.00134 0.9959	-0.25374 0.3430 16
90	1.00000 0.00000 15	-0.07207 0.7763	-0.29363 0.2697 16	-0.16819 0.4695 17	-0.24672 0.3570 16
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VARIABLE	z	HEAN	STD DEV	NUS .	HUMINIH	HAXINU
90	8	79.16666667	21.25268149	1425.0000000	146.0000000	105.000000
Ħ	91	-89.5000000	90.20124559	-1611.0000000	-280,0000000	5.000000
COND	11	329.11764706	46.03786804	5595.0000000	260.0000000	400.000000
CHL	11	32.58823529	4.96309913	554.0000000	25.0000000	45.000000
•	11	0.20470566	0.14688931	3.4800000	0.0300000	0.550000

CORRELATION COEFFICIENTS / PROB > IR| UNDER NO:RHO=D / NUMBER OF OBSERVATIONS

•	0.14122 0.5888	-0.48480 0.0486	0.19012 0.4649	-0.02204 0.9331	0.0000
Ŗ	0.10469 0.6893	0.40332 0.1084 17	0.23629 0.3612 17	1.00000 0.0000 17	-0.02204 0.9331
H CÓND	-0.31739 0.2145 71	-0.26538 0.3033	0.0000	0.23629 0.3612	0.19012 0.4649 17
Ŧ	0.09275 0.7143	1.00000 0.0000 18	-0.26538 0.3033	0.40332 0.1084 17	-0.48480 0.0486
90	1.00000 0.0000 81	0.09275 0.7143 81	-0.31739 0.2145 17	0.10469 0.6893	0.14122 0.5666 17
	40	Ŧ	COND	¥	•

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APPENDIX C

PEARSON PRODUCT MOMENT CORRELATIONS OF OBSERVATIONS CLASSIFIED ACCORDING TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY

		•7	TATISTICAL	STATISTICAL ANALYSIS SYSTEM Sthred	4 17:09 FRIDAY, OCTOBER 30, 1981	TOBER 30, 1981
VARIABLE		HEAN	STD DEV	MNS	HUMININ	MAXIM
6	59	78,0000000	22.01919260	1602,0000000	35.000000	123.000000
Ŧ	75	-36.48148148	82.72385085	-1970.0000000	-360.0000000	70.00000
COND	25	321.92982456	50.40686590	18350.0000000	250.0000000	540.000000
CIIL	57	29.85964912	3.28110715	1702.0000000	25.0000000	40.000000
•	5	0.28927273	0.39676868	15.9100000	•	2.500000
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<u>-</u>	1.00000	-0.22424 0.1031	0.16919 0.2083	0.20429 0.1274	0.37
Ŧ	-0.22424 0.1031	1.00000 0.0000 54	0.34441 0.0116	0.02999 0.8312 53	-0.00
QNO	0.16919 0.2083	0.34441 0.0116 53	1.00000 0.0000 57	0.43355 0.0008 57	0.050
ž	0.20429 0.1274 57	0.02999 0.8312	0.43355 0.0008 57	1.00000 0.0000 57	0.030
	0.13787 0.3155 55	-0.00475 0.9734 52	0.05026	0.03011 0.8273	0.00

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		5 T A T I S	TICAL ANAL STNK=1	STATISTICAL ANALYSIS SYSTEM 17:09 FRIDAV, OCTOBER 30, 1981 STATISTICAL ANALYSIS SYSTEM	17:09 FRIDAY, 0	CTOBER 30, 1981
VARIABLE	-	HEAN	STD DEV	NUS	NUMININ	INNIXYN
40	36	74.6666667	19.18332609	2688.0000000	116.0000000	113.000000
ни	36	-76.7500000	86.34494774	-2763.0000000	-280.0000000	7.000000
COND		340.454545	66.79450338	11235.0000000	260.0000000	620.0000000
CHL	34	31.41176471	5.00979433	1068.0000000	20.0000000	45.000000
٩.		0.214242	0.19854203	7.0700000	•	1.000000
		CORRELATION COEFFICIENTS / PROB > IR UNDER HO:RHO=D / NUMBER OF OBSERVATIONS	/ PROB > IR} UNDER	HO:RHO=D / NUMBER OF OBS	ERVATIONS	

4 2 2 2

•	-0.07081 0.6954	-0.32440 0.0655	0.02777	0.20827 0.2448	1.00000 0.0000 33
CHL	0.04008 0.8220 34	0.17022 0.3358 34	-0.17517 0.3295 33	1.00000 0.0000 34	0.20827 0.2448 33
COND	-0.30350 0.0860 33	0.16990 0.3445 33	1.00000 0.0000 33	-0.17517 0.3295 33	0.02777
Ŧ	-0.01190 0.9451 36	1.00000 0.0000 36	0.16990 0.3445 33	0.17022 0.3358	-0.32440 0.0655 33
-0	1.00000 0.0000 36	-0.01190 0.9451 36	-0,30350 0,0860 33	0.04008 0.8220 34	-0.07081 0.6954 33
	6	Ŧ	COND	CHL	•

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APPENDIX D

ANALYSIS OF OBSERVATIONS CLASSIFIED ACCORDING TO VALUES ON BOTH INDEPENDENT VARIABLES USING A GENERAL LINEAR MODELS PROGRAM

STATISTICAL ANALYSIS SYSTEM 18:20 FRIDAY, OCTOBER 23, 1981

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GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

00100	LLILLU	TALULU
STNK	2	0 1
DIST	2	10 20

NUMBER OF OBSERVATIONS IN DATA SET = 95

GROUP	OBS	DEPENDENT VARIABLES
1	90	HH
2	90	COND
3	91	CHL
4	88	,

NOTE: VARIABLES IN EACH GROUP ARE CONSISTENT WITH RESPECT TO THE PRESENCE OF ABSENCE OF MISSING VALUES.

87

STATISTICAL ANALYSIS SYSTEM 18:20 FRIDAY, OCTOBER 23, 1981 7

GENERAL LINEAR MODELS PROCEDURE

ILE: HH							
DF	SUM OF SQUARES	MEAN S	QUARE	F VALUE	PR > F	R-SQUARE	c.v.
3	72366.99259259	24122.330	86420	3.54	0.0179	0.109870	157.0049
86	586290.79629630	6817.334	84065		STD DEV		HH MEAN
89	658657.78888889				82.56715352	-5	2.58888889
DF	TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS	F VALUE	PR > F
1.	35025.55740741 34535.21111111 2806.22407407	5.14 5.07 0.41	0.0259 0.0270 0.5228	;	35025.55740741 29407.82407407 2806.22407407	5.14 4.31 0.41	0.0259 0.0408 0.5228
	DF 3 86 89	DF SUH OF SQUARES 3 72366.99259259 86 586290.79629630 89 658657.7888889 DF TYPE I SS 1 35025.55700711 1 35035.2111111	DF SUM OF SQUARES MEAN S 3 72366.99259259 24122.30 86 566250.786286389 DF TYPE ISS F VALUE 1 330625.55700711 1 35055.5700711	DF SUM OF SQUARES HEAN SQUARE 3 72366.99259259 24122.33066420 86 586290.79629630 6817.33444065 89 655657.78686889 DF TYPE I SS F VALUE PR > F 1 34055.55780711 5.14 0.0259 1 34055.55780711 5.14 0.0259	DF SUM OF SQUARES HEAN SQUARE F VALUE 3 72366.0929529 20122.30060x20 3.54 86 566290.79629630 6817.33484065 3 89 65657.78888889 5 5 DF TYPE I SS F VALUE FR > F DF 1 35035.537107141 5.14 0.0250 1	DF SUM OF SQUARES HEAN SQUARE F VALUE PR > F 3 72366.09259259 24122.30060420 3.54 0.0179 86 566290.79629630 6817.334840655 STD DEV 89 656657.78688689 62.56715352 DF TYPE I SS F VALUE PR > F 1 35025.55710741 5.14 0.0259 1 3 3553.25171011 5.14 0.0259 1 35025.55710741	DF SUM OF SQUARES HEAN SQUARE F VALUE PR > F R-SQUARE 3 72366.9829529 24122.33086420 3.54 0.0179 0.109870 86 586230.79529530 6817.33484065 STD DEV 89 658657.78888899 82.3511332 -5 DF TYPE IS F VALUE PR > F DF TYPE IV S5 F VALUE 1 35025.5570011 5.14 0.0259 1 35027.5807011 5.14 1 35032.5571011 5.14 0.0259 1 35027.5807011 5.14

STATISTICAL ANALYSIS SYSTEN 18:20 FRIDAY, OCTOBER 23, 1981 8

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GENERAL LINEAR MODELS PROCEDURE

DEPENDENT VAR	TABLE: COND								
SOURCE	OF	SUM OF SQUARES	HEAN SO	UARE	F VALUE	PR > F	R-SQUARE	c v.	
MODEL	3	13386.11843588	4462.0394	7863	1.38	0.2546	0.045807	17.3221	
ERROR	86	278841.93711968	3242.3481	0604		STO DEV		COND MEAN	
CORRECTED TOT	AL 89	292228.05555556				56.94162016	3	28.72222222	
SOURCE	DF	· TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS	F VALUE	PR > F	
SINK Dist Stnk®dist	1	7172.15443913 60.09733392 6153.866666283	2.21 0.02 1.90	0.1406 0.8920 0.1719	ļ	7368.39148553 807.25673375 6153.86666283	2.27 0.25 1.90	0.1353 0.6191 0.1719	

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STATISTICAL ANALYSIS SYSTEN 18:20 FRIDAY, OCTOBER 23, 1981

GENERAL LINEAR MODELS PROCEDURE

DEPENDENT VARIABLE: C	HL.							
SOURCE	DF	SUM OF SQUARES	MEAN S	QUARE	F VALUE	PR > F	R-SQUARE	c.v.
MODEL	3	98.42460937	32.808	20312	2.06	0.1095	0.066395	13.1029
ERROR	87	1383.99297305	15.907	96521		STD DEV		CHL MEAN
CORRECTED TOTAL	90	1482.41758242				3.98847906		30.43956044
SOURCE	DF -	TYPE I SS	F VALUE	PR > F	DF	TYPE IN SS	F VALUE	PR > F
STNK DIST STNK#DIST	1	51.30509532 19.25899013 27.86052392	3.23 1.21 1.75	0.0760 0.2742 0.1892	ł	51.26135421 31.13086715 27.86052392	3.22 1.96 1.75	0.0761 0.1654 0.1892

STATISTICAL ANALYSIS SYSTEM 18:20 FRIDAY, OCTOBER 23, 1981 10

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GENERAL LINEAR HODELS PROCEDURE

DEPENDENT VARIABLE:	P							
SOURCE	DF	SUM OF SQUARES	MEAN S	QUARE	F VALUE	PR > F	R-SQUARE	C.V.
MODEL	3	0.24421604	0.08140535		0.71	0.5523	0.024722	129.6888
ERROR	84	9.63427032	0.11469369		STD DEV			P HEAN
CORRECTED TOTAL	87	9.87848636				0.33866457		0.26113636
SOURCE	DF -	TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS	F VALUE	PR > F
STNK Dist Stnk*dist	1.	0.11610939 0.05996307 0.06814358	1.01 0.52 0.59	0.3172 0.4717 0.4430	1	0.11245420 0.02949724 0.06814358	0.98 0.26 0.59	0.3249 0.6134 0.4430

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68

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APPENDIX E

ANALYSIS OF OBSERVATIONS CLASSIFIED ACCORDING TO THE PRESENCE OR ABSENCE OF A SEPTIC TANK ONLY USING A GENERAL LINEAR MODELS PROGRAM

STATISTICAL ANALYSIS SYSTEM 11:09 FRIDAY, OCTOBER 30, 1981

GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION CLASS LEVEL INFORMATION

STNK Z 01

IUMBER OF OBSERVATIONS IN DATA SET = 95

UP OBS DEPENDENT VARIABLES

HH 06

90 CONI

91 CHL

•

NOTE: VARIABLES IN EACH GROUP ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES.

STATISTICAL ANALYSIS SYSTEM 17:09 FRIDAY, OCTOBER 30, 1981 4

GENERAL LINEAR MODELS PROCEDURE

DEPENDENT VARIABLE: HH								
SOURCE	DF	SUH OF SQUARES	MEAN SQU	ARE	F VALUE	PR > F	R-SQUARE	c v.
MODEL	1	35025.55740741	35025.55740	741	4.94	0.0288	0.053177	160.0776
ERROR	88	623632.23148148	7086.72990	320		STD DEV		HH MEAN
CORRECTED TOTAL	89	658657.78888889				84.18271737		-52.58888889
SOURCE	DF	TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS	F VALUE	Pit > I
STNK	1	35025.55740741	4.94	0.0288	1	35025.55740741	4.94	0.0288

STATISTICAL ANALYSIS SYSTEM 17:09 FRIDAY, OCTOBER 30, 1981 5

GENERAL LINEAR HODELS PROCEDURE

DEPENDENT VARIABLE:	COND							
SOURCE	DF	SUM OF SQUARES	MEAN SQUAR	E FV	ALUE	PR > F	R-SQUARE	c.v.
MODEL	1	7172.15443913	7172.1544391	3 :	2.21	0.1403	0.024543	17.3139
ERROR	88	285055.90111643	3239.2716036	0		STD DEV		COND MEAN
CORRECTED TOTAL	89	292228.05555556				56.91459921		328.7222222
SOURCE	DF	. TYPE I SS	F VALUE	R > F	DF	TYPE IV SS	F VALUE	PR > F
STNK	1	7172.15443913	2.21	. 1403	1	7172.15443913	2.21	0.1403

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GENERAL LINEAR MODELS PROCEDURE

DEPLNDENT VARIABLE: CHL	сн							
SOURCE	Ъ	SUM OF SQUARES	HEAN SQI	NARE	F VALUE	PR > F	R-SQUARE	с. v.
HODEL	-	51.30509532	51.30509532	9532	3.19	0.0775		13.1736
ERROR	89	1431.11248710	16.07991	1559		STD DEV		CHL MEAP
CORRECTED TOTAL	8	1482.41758242				4,00997701	-	0.43956044
SOURCE	D	· TYPE I SS	F VALUE	PR > F	Q	TYPE IV SS	F VALUE	PR > F
STHK	-	51.30509532	3.19	0.0775	-	51.30509532		0.0775
	•							

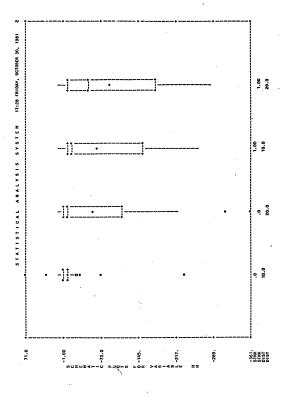
STATISTICAL ANALYSIS SYSTEM 17:09 FRIDAY, OCTOBER 30, 1981

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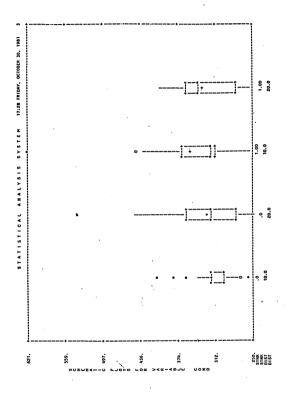
<u>د</u>. PR > F 129.0212 7 NEAP 0.26113636 0.3147 1.02 F VALUE **R-SQUARE** 0.011754 TYPE IV 55 0.11610939 STD DEV PR > F 0.3147 0.33692137 F VALUE 1.02 h • GENERAL LINEAR MODELS PROCEDURE 0.3147 PR > F MEAN SQUARE 0.11610939 0.11351601 1.02 F VALUE SUM OF SQUARES 0.11610939 9.76237697 TYPE I SS 0.11610939 9.87848636 5 8 DEPENDENT VARIABLE: P CORRECTED TOTAL SOURCE SOURCE ERROR HODEL XNI

APPENDIX F

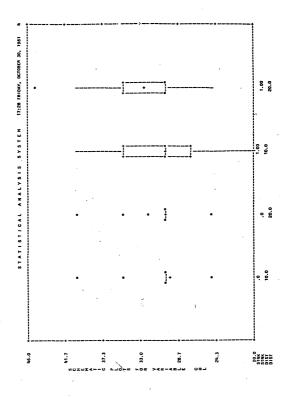
SCHEMATIC PLOTS FOR THE DEPENDENT VARIABLES



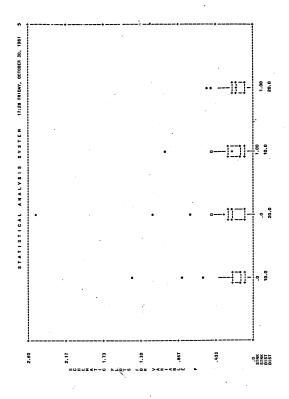
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APPENDIX G

EXAMINATION FOR SKEWNESS USING THE UNIVARIATE PROCEDURE

11:54 SATURDAY, NOVEMBER 14, 1981 EXTREMES 1999 EXMINATION OF DISTRIBUTIONS FOR SKEWNESS UNIVARIATE

VAR1ABLE=HH	H							
	HOHE	IOHENTS			QUANT I LES (DEF=4)	DEF=4)		
N HEAN STD DEV Skenness USS CV T:MEAN=0	-52.5889 86.027 -1.69155 -163.584	SUH MGTS SUH MGTS VARIANCE VARIANCE KURTOSIS CSS STD HEAN PROB-[T]	-4733 -4733 7400.65 2.13861 658658 9.06804 0.0001	1004 HAX 754 03 2554 HED 2554 01 034 HED 034 HED 034 HED	71.25 1360 1360	6666- 888888	5.89999 1.89999 -246.75 -360	9
				HODE MISSING VALUE COUNT/NOBS	ALUE D OUNT 5.26 NOBS 5.26			

225 HIGHEST

4

ι

EXTREMES

22022862

SCOOL ST

ARIANCE ARIANCE AURTOSIS SID HEAN PROB-111

HIGTS

53

822

82

HEAN

200

100E

VALUE

MISSING

COUNT

EXAMINATION OF DISTRIBUTIONS FOR SKEWNESS

TI:54 SATURDAY, NOVEMBER 14, 1981

UNIVARIATE

UANTILES(DEF=4) 29292

IOMENTS VARIABLE=COND

11:54 SATURDAY, NOVENBER 14, 1981 EXAMIMATION OF DISTRIBUTIONS FOR SKEWNESS UNIVARIATE

EXTREMES

PROPERTY.

222 HE 222

908258

RANGE Q3-Q1 HODE

VALUE COUNT MOBS

DNISSING X COUNT

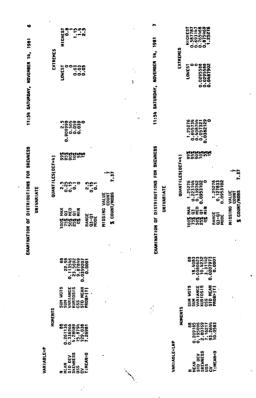
QUANTILES(DEF=4)

5.8861

0515 5 **DHENTS** VARIABLE=LNCOND

HIGHEST 11:54 SATURDAY, NOVEMBER 14, 1981 11:54 SATURDAY, NOVEMBER 14, 1981 EXTREMES EXTREMES 282220 2822202 3322008 EXAMINATION OF DISTRIBUTIONS FOR SKEWNESS EXAMIMATION OF DISTRIBUTIONS FOR SKEWNESS SUCCESSION. (DEF=4) HHHHH QUANTILES(DEF=4) 12 32222 500 90666 0100 UNIVARIATE UNIVARIATE MISSING VALUE COUNT COUNT/NOBS ¥896 MISSING X COUNT ¥8 11SSII RANCE 03-01 HODE 1002 SUH MGTS SUH VARIANCE KURTOSIS CSS STD MEAN PROB>IT Ξ÷ SUH MOTS AR I ANCE SS STD HEAN PROB-|T| TOHENTS HOMENTS VARIABLE=LNCIIL VARIABLE=CHL HFAN= I HIN STO DEV SKENNE AN V HEAN

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APPENDIX H

FINAL ANALYSIS OF HYDRAULIC HEAD USING THE GENERAL LINEAR MODELS PROCEDURE

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104

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SECOND PRINT OUT--HAS DW AND DP--DDES NOT HAVE DIST THIS SEEMS TO BE THE BEST MODEL FOR HH

13:23 SATURDAY, NOVEMBER 14, 1981

GENERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

CLASS	LEVELS	VALUES
GL.	3	
SOIL	6	a b o1 o2 o3 o4
PHBL	3	123

NUMBER OF OBSERVATIONS IN DATA SET = 95

NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OF ABSENCE OF MISSING VALUES. HOWEVER, ONLY 87 Observations in data set can be used in this analysis.

2 x		139	Ŧ	-51.356	-	000000												
13:23 SATURDAY, MOVEMBER 14,	R-SQUARE	0.393400		Ť	F VALUE	20 20 20 20 20 20 20 20 20 20 20 20 20 2						E USED.	8	0.9315 0.7955 0.1367 0.0069 0.0401	E USED.			
	PR > F	0.0001	STD DEV	71.77036841	TYPE IV SS	39901.07561094 54158.81373570 99144.44790641 27479.59268100 222661.01277721 22088.00577280				PRO8 > 111 HO: LSHEAN(1)=LSHEAN(J) 1/J 1 2 3	0.8897 0.0138	ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.	LSHEAN(J) 5	0.0395 0.1725 0.3297 0.6228 0.6748 0.6822 0.2804 0.2804 0.20059 0.0401	PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.	PROB > 171 HO: LSHEAN(1)=LSNEAN(J) 1/J 1 2 3	0.5130	
THIS SEEKS TO BE THE DEST MODEL FOR HH General Linear Models procedure	F VALUE	4.00				~~~~~	STD ERROR OF Estimate	9.81262740 0.37303384 0.38572436		 ITI H0: LSME 1 2 	0.0110 0.0110 0.0138	PRE-PLANNED CO	11 HO: LSNEAN(I)=LSHEAN(J)	0.8366 0.2235 0.5172 0.5172 0.3297 0.6748 0.6228 0.6822 0.6228 0.1367	PRE-PLANNED CO	> 171 HO: LSM	0.0007 0.0007 0.5130	
S SEEMS TO BE THE BEST MODEL FOI Gemeral Linear Models Procedure	MEAN SQUARE	051561	578156		PR > F	0.0038 0.1924 0.0004 0.1082 0.0525 0.0525	STC		LEAST SQUARES MEANS		0.0260 1 0.0003 2 0.0	CIATED WITH	111 ~ 8084	2 0.8366 9 0.0395 5 0.1725 5 0.9315	CIATED WITH		0.0331 2 0	
MAL LINEAR	MEAN	20600.25051561	5150.98578156		F VALUE	6.03 9.75 3.88 4.29	PR > T	0.0237 0.0419 0.0411	LEAST SQU	PROB > IT HO:LSHEAN=0		ITIES ASSO	PROB > [T] H0:LSMEAN=0	0.0019 0.1852 0.3664 0.6013 0.0182 0.001	ITIES ASSO	R PROB > 11		
THIS SET	SUM OF SQUARES	0618733	4783566	5402299	TYPE I SS	62104.94037356 39196.56209504 90188.22659953 13615.26705871 20009.92429768 22088.08577280	T FOR HO: PARAMETER=0	-2.31		STD ERR LSMEAN	23.248604 34.475071 19.278604	WLY PROBABIL	STD ERR PRO	36.601938 75.726155 48.151513 35.716058 26.040672 26.040672	WLY PROBABIL	STD ERR LSMEAN	33.486766 20.106438 22.899093	
	SUM OF	247203.00618733	381172.94783566	628375.95402299	71	62104.9 39196.5 90188.2 13615.2 20009.9	PAR			HH LSHEAN	-52.810557 -130.571157 -49.476125				LEVEL, 0	HH LSMEAN	-159.003278 -43.665534 -30.191028	
_	DE	12	2	98	10	NON	ESTIMATE	-22.66445491 -0.77247085 -0.80187072		Ъ	52 52 52	MOTE: TO ENSURE OVERALL PROTECTION LEVEL,	LSHEAN	-118.040351 -101.264237 -43.757861 -18.741664 -18.741664 -62.893036	L PROTECTION	JBH4	- 0 m	
RIABLE: HH				١٧٢					,			RE OVERAL	SOIL	**2 <u>88</u> 9	RE OVERALI			
DEFCNDENT VARIABLE: HH	SOURCE	HODEL	ERROR	CORRECTED TOTAL	SOURCE	SSIL SSIL DBA DBA	PARAMETER	DDA LINEAR Dy Linear DP Linear				E: TO ENSU			NOTE: TO ENSURE OVERALL			

, 1981

C.V. 9.7498 H HEAN 632184 PR > F

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APPENDIX I

FINAL ANALYSIS OF SPECIFIC CONDUCTIVITY, CHLORIDE CONTENT, AND PHOSPHATE CONTENT USING THE GENERAL LINEAR MODELS PROCEDURE

14:01 SATURDAY, NOVEMBER 14, 1981

FULL MODEL FOR LOG OF CONDUCTIVITY General Linear Models Procedure Class level information

CLASS LEVEL INFORMATION CLASS LEVELS VALUES

3 cmo 6 = bolo2 o3 o4

UMBER OF OBSERVATIONS IN DATA SET = 95

5 NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOMEVER, ONLY Observations in data set can be used in this analysis.

DEFENDENT VARIABLE: LNCOND	LNCOND							
SOURCE	DF	SUM OF SQUARES	MEAN SQUARE	QUARE	F VALUE	PR > F	R-SQUARE	c. v.
HODEL	15	0.41659720	0.02777315	21577	1.17	0.3151	0.207772	2.6591
ERROR	67	1.58847276	0.02370855	70855		STD DEV	-	NCOND MEAN
CORRECTED TOTAL	20	2.00506996				0.15397581		5.79053374
SOURCE	10	TYPE 1 SS	F VALUE	PR > F	ų	TYPE IV SS	F VALUE	PR > F
SOIL	~~~	0.05843242 0.05131437	1.23	0.2981 0.8255	~~~	0.08529341 0.03299559		0.1734 0.9227
DIST	·	0.00046802	0.02	0.8887		0.01004486		0.5173
HH DBA		0.06685864 0.02841662	2.92	0.0978 0.2775		0.00710891	0.9	0.0788
M		0.06093163	2.51	910		0.06093163	2.57	0.1136
PARAHETER	ESTIMATE	T FOR HO: PARAMETER=0	PR > ITI	15	STL ERROR OF			
OBA LINEAR DV LINEAR DP LINEAR HH LINEAR	-0.01228158 0.00156221 0.00046345 0.000468345	-0.55 1.60 1.10	0.5858 0.1136 0.6249 0.0788		0.02242875 0.00097448 0.00094253 0.00026255			

108

FULL MODEL FOR LOG OF CONDUCTIVITY

14:01 SATURDAY, MOVEMBER 14, 1981

GENERAL LINEAR MODELS PROCEDURE

LEAST SQUARES MEANS

+ 171 HO: LSMEAN(1)=LSMEAN(J) 1 2 3	0.9834 0.0717
Ξ-	÷
PR08	
-	_
FROD > 111 HO:LSHEAN=0	0.0001
STD ERR LSHEAN	0.05167406
LINCOND	5.69286462
ಠ	0

BE USED. NOTE: TO ENSURE OVERALL

9	0.7385 0.4348 0.4329 0.6815	E USED.
~	0.59618	SHOULD B
HEAN(J)	.4212 .3818 .3818 .5811	ARISONS
3 3		INED COM
HO: LSM 2	0.5606	PRE-PLAI
PROB > 11 HO: LSMEAN(I)=LSMEAN(J) 1/J 1 2 3 4	0.5606 0.6015 0.9818 0.7385	TED WITH
₹≥	-00300	ŝ
PROB > 111 H0:LSMEAH=D	000.00000000000000000000000000000000000	HOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED.
STD ERR 1. SHEAN	0.08056614	" ONLY PROB
	000000	EVEL
LNCOND	5.744143409 5.64046158 5.67300035 5.79242649 5.770242649 5.770242230	PROTECTION L
SOIL	• 16 8 8 8 4 •	OVERALL
		O ENSURE
1		1
		NOTE

PROB > 171 HO: LSHEAN(I)=LSHEAN(J) 1/J 1 2 3 PROB > 111 H0:LSHEAN=0 STD ERR LSHEAN LNCOND PMBL

	5	
	SHOULD	
0.0001 2 0.1455 0.7959 0.7959	COMPARI SONS	PROB > T HO: LSMEAN1=LSMEAN2
0.795	LANNED	PROB >
0.1455	TH PRE-P	PROB > 111 HO:LSMEAN=0
· (1 m)	ž	800
0.000	ASSOCIATE	STD ERR PI
0.05042015	PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD	LINCOND SI
5.75648829	ECTION LEVEL,	DIST
	ROT.	

0231 0

E USED. MOTE: TO ENSURE OVERALL PF

PROB > 1T1 HO: LSNEAN1=LSNEAN2	0.5173	PROB > 11 HO: LSHEAN1=LSHEAN2	0.0277
PROB > 11 HO:LSMEAN=0	0.0001	PROB > 111 HO:LSMEAW=0	0.00010
STD ERR LSMEAN	0.04797368	STD ERR LSMEAN	0.05354731
LNCOND	5.74088800	LNCOND	5.1771814
DIST	n 10	STNK	o-

REDUCED MODEL FOR LOG OF CONDUCTIVITY--INI AND SINK ONLY MODEL INCLUDES TEST FOR INTERACTION BETHEEN HH AND SINK

4 14:01 SATURDAY, NOVEMBER 14, 1981

GEMERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

VALUES LEVELS CLASS

5 ~ STNK NUMBER OF OBSERVATIONS IN DATA SET = 95

8 NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOMEVEM, OMLY Observations in data set can be used in this amalysis.

	DEPENDENT VARIABLE: LNCOND	LNCOND							
	Source	'n	SUM OF SQUARES	HEAN S	QUARE	F VALUE	PR > F	R-SQUARE	<u>,</u>
- 1	, HODEL		0.22103629	0.073	0.07367876	3.29	0.0246	0.107320	2.586
۰.	ERROR	8 2	1.63857037	0.022	42159		STD DEV		LNCOND MEA
	CORRECTED TOTAL	5	2.05960666				0.14973840		5.7885820
	SOURCE	. 10	TYPE I SS	F VALUE	PR > F	Dr .	TYPE IV SS	F VALUE	PR •
	STNK HII HII=STNK		0.03663481 0.16130638 0.02309510	1.63	0.2048		0.02028803 0.13468570 0.02309510	0.90	0.314
	PARAMETER	ESTIMATE	T FOR HO: Parameter=0	PR > 11	STD E	STD ERROR OF ESTIMATE			
	HH LINEAR	0.00047767	2.45	0.0164	•	0.00019489			

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335

PROB > [T] HO:LSHEAN=0 0.0001 LEAST SQUARES MEANS STD ERR LSHEAN 0.02100802 0.02734543 LNCOND 5.76069992 STNK

PROB > 1T HO: LSHEAN1=LSMEAN2 0.0770 14:01 SATURDAY, NOVEMBER 14, 1981

FULL MODEL FOR LOG OF CHLORIDE GENERAL LINEAR MODELS PROCEDUR CLASS LEVEL INFORMATION VALUES • LEVELS CLASS 2

b o1 o2 o3 o4 10 151 STHK UNBER OF OBSERVATIONS IN DATA SET = 95

2 NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSIMG VALUES, HOMEVEN, ONLY Observations in data set can be used in this Amalysis.

DEPENDENT VARIABLE: LNCHL	1CHL							•
SOURCE	Dr	SUM OF SQUARES	MEAN SQUARE	RE	F VALUE	PR > F	R-SQUARE	C.V.
MODEL	15	0.36641465	0.02442764	191	1.53	0.1198	0.252024	3.7062
ERROR	68	1.08747393	0.01599226	26		STO DEV		LNCHL NEAN
CORRECTED TOTAL	83	1.45388658				0.12646052		3.41211025
SOURCE	DŁ	TYPE I SS	F VALUE	PR > F	DE .	TYPE IV SS	F VALUE	PR > F
Solt.	Q	0.07914713	2.47	0.0918	N 17	0.00763736	0.25	0.7834
TIRE	\Q P	0.06580013	200	0.1357		0.07615292	5.3	0.1001
STNK		0.03520335	2.20	0.1425		0.01494383	0.93	0.3371
VBD		0.00602627	0.36	0.6020		0.00347926	0.22	0.6424
10		0.05699377	3.56	0.0633		0.02647777 0.06611962	1.66	0.2026
PARAHETER	ESTIMATE	T FOR HD: Parameter=0	PR > [T].	STO	STD ERROR OF Estimate			
DBA LINEAR DV LINEAR DP LINEAR HH LINEAR	0.00313597 -0.00162635 0.000199526 0.00010050	-2.03	0.8652 0.0159 0.2026 0.6424		0.01839911 0.00079984 0.00077348 0.000277348			

111

14:01 SATURDAY, NOVEMBER 14, 1981

JEMERAL LINEAR MODELS PROCEDURE LEAST SQUARES MEANS

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED

PROB > [T] HO: LSHEAN(!)=LSHEAN(J)

STD ERR LSMEAN 0.09237738 0.06528690 0.04818223 0.02921968 0.06629866

SOL

FULL HODEL FOR LOC OF CHLORIDE

PROB > 171 HO: LSHEAH(I)=LSHEAN(J) 1/J 1 2 3

PROB > 111 HO:LSHEAN=0

STD ERR LSHEAN

LNCHL

H

2 0.6274 0.6274 0.5450 2 0.6274 0.9839 3 0.5450 0.9839 .

0.0001

0.04249541 0.06470447 0.03516867

3.44599750 3.47311118 3.47438788

112

ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD 2 0.1268 3 0.7123 ONLY PROBABILITIES PROTECTION LEVEL, NOTE: TO ENSURE OVERALL

PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED

PROB > [T] HO: LSMEAN(I)=LSMEAN(J) 1/J 1 2 3

PROB > 111 HO:LSHEAN=0

STD ERR LSMEAN

LNCHL

R

NOTE: TO ENSURE OVERALL

0.1268 0.7123

0.0467

0.0001

0.06340807 0.03650618 0.04105315

3.50383082 3.40751445 3.46215129

0.1421 0.5775 0.6468 0.9849 0.3637

. 4194 0.7319

0.2226 0.3942 0.5400 0.3637

0.8643

PROB > 111 HO:LSHEAN=0

3.51135602 3.53637960 3.47289126 3.40427498 3.40427498 3.47124613 3.39084214 LINCHL

PROB > 11 HO: LSMEAN1=LSHEAN2	0.1639	PROB > 171 HO: LSMEAN1=LSHEAN2	0.3371
PROB > T HO:LSMEAN=0	0.0001	PROB > [T] HO:LSMEAN=0	0.001
STD ERR LSHEAN	0.03932294 0.04267950	STD ERR LSMEAN	0.04399166
LINCHL	3.44020745 3.48879025	LINGHL	3.461723775
DIST	n 10	STNK	o-

REDUCED HODEL FOR LOG OF CHLORIDES--HII AND STNK ONLY MODEL INCLUDES TEST FOR INTERACTION BETWEEN HH AND STNK

10 14:01 SATURDAY, NOVEMBER 14, 1981

GEMERAL LINEAR MODELS PROCEDURE

CLASS LEVEL INFORMATION

VALUES LEVELS CLASS

-• STAK NUMBER OF OBSERVATIONS IN DATA SET = 95

5 NOTE: ALL DÉPENDENT VASIABLES ARE CONSISTENT MITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOMEVER, ONLY Observations in data set cam be used in this Amalysis.

Mont. 1 0.01953041 0.9195041 0.9185004 0.918500 0.018300 0.010300 RHMA 81 1.40054221 0.01893044 91 97 97 91 </th <th>SOURCE</th> <th>or</th> <th>SUM OF SQUARES</th> <th>MEAN SQUARE</th> <th></th> <th>F VALUE</th> <th>PR > F</th> <th>R-SQUARE</th> <th></th>	SOURCE	or	SUM OF SQUARES	MEAN SQUARE		F VALUE	PR > F	R-SQUARE	
0 1.40634821 0.6166444 510 DV 1 6 1.40634821 0.6166444 9.10 DV 1 16 1.40634821 0.416 M×1 9.1 0.4101192 1 1 0.3030001 1.71 0.1411792 1 0 0.0331963 0.14 1	HODEL	8	0.04785133	0.0159	5044	0.94	0.4262	0.032905	
66 1,49423334 0,11011032 1 0r 1,7071 35 r VAUK m > r 0 1 0 0r 1,0023344 0,51 0,1101 0 0 1 0 0 1 0	ERROR	63	1.40638221	0.0169	9644		STD DEV		LNCHL 1
DT VTC 1 is 5 F VALUE m > F DF TVTC 1 vs 5 F V 1 0.000010000 0.231 0.0000 0.00001000 0.00001000 0.00001000 1 0.00001000 0.331 0.0001 0.0001000 0.0001000 0.0001000 0.0001000 0.0001 0.0001000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.00010000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.000100000 0.000100000 0.00010000 0.00010000 0.0001000 0.00010000 0.00010000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.00010000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.00010000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.00010000 0.00010000 0.00010000 0.00010000 0.0001000 0.00010000 0.000100000 0.00010000000000 0.000000000	CORRECTED TOTAL	98.	1.45423354				0.13017052		3.4117
1 0.000000 1.17 0.167 1 0.0004569 1 0.0001599 0.53 0.043 1 0.044469 1511MATE 0.0001599 0.43 0.167 1 0.0001595 0.00015999 0.43 0.403 0.00016844 0.00016844 0.00016844 11441 0.00016844 0.00016844 0.00016844 0.00016844 0.00016844 11441 0.01001644 0.00016844 0.00016844 0.00016844 0.00016844	SOURCE	5	TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS		R
ESTIMATE TOGAIN: PA> 111 STERIMATE 0.00013999 0.433 0.4043 0.000148 0.00013999 0.443 0.4043 0.000148 0.4443 0.4443 0.44445 0.44445 1.1445 0.44445 0.44445	STNK HII HH®STNK		0.03003043 0.00864444	1.77	0.1867 0.4771 0.4639		0.01170479 0.001170479 0.00917645		
0.00013999 0.83 0.403 0.00014 1.1431 000045 News 1.1431 000045 News 1.1431 000045 News 1.1431 000045 News	PARAHETER	ESTIMATE	T FOR HO: PARAMETER=0	PR > T	STO ERI ESTI	KOR OF MATE			
LEAST SQUARES HEANS LEAST SQUARES HEANS STNK LIVEN LSTREAM POLISECANED	HH LINEAR	0.00013999	0.63	0.4083	0.00	16843			
LEAST SOUARES HEANS LNCHL STD ERR PROB > 171 LSNEAN LSNEAN=D		[•	ı	
LNCHL STD ERR PROB > [T] LSWEAN LSHEAN HO:LSMEAN=O				LEAST SQUA	RES MEANS				
			-	STD ERR LSHEAN			IT HO: -LSMEAN2		

0.1305

0.0001

0.01824054

3.39658611

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C. V. 8154 NEAN 73395

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14:01 SATURDAY,	NOVEMBER 14	. 1981 13
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GENER/	L LINEAR H	ODELS PROCEDURE
CI	ASS LEVEL	INFORMATION
CLASS	LEVELS	VALUES
GL.	3	
SOIL	6	a b o1 o2 o3 o4
PMBL	3	1 2 3
DIST	2	36
STNK	2	0 1

FULL MODEL FOR LOG OF PHOSPHATES

NUMBER OF OBSERVATIONS IN DATA SET = 95

NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOWEVER, ONLY Observations in data set can be used in this analysis.

DEPENDENT VARIABLE	E: LNP							
SOURCE	DF	SUM OF SQUARES	MEAN SO	UARE	F VALUE	PR > F	R-SQUARE	c.v.
MODEL	15	0.72250670	0.048	6711	1.27	. 0.2450	0.224221	90.3173
ERROR	66	2.49979115	0.0376	37562		STO DEV		LNP HEAN
CORRECTED TOTAL	81	3.22229785				0.19461661		0.21548100
SOURCE	DF ·	TYPE I SS	F VALUE	PR > F	DF	TYPE IV SS	F VALUE	PR > F
GL SOIL DIST STMK HIH DBA DP DW	252111111111111111111111111111111111111	0.28304683 0.19696296 0.14%27385 0.03042560 0.00002094 0.02911663 0.00180485 0.03630822 0.00024683	3.74 1.04 0.80 0.00 0.78 0.05 0.96 0.01	0.0290 0.4022 0.1570 0.3734 0.9813 0.3814 0.8279 0.3311 0.9359	252111111111111111111111111111111111111	0.28374077 0.11910154 0.14537895 0.00327691 0.00187160 0.01950469 0.00037834 0.03231544 0.00024683	3.75 0.63 1.92 0.05 0.5 0.51 0.01 0.85 9.01	0.0288 0.6806 0.1548 0.7696 0.8246 0.4755 0.9207 0.9207 0.3590 0.9359
PARAMETER	ESTIMATE	T FOR HO: PARAMETER=0	PR > ITI	5	D ERROR OF ESTIMATE			
DBA LINEAR DW LINEAR DP LINEAR HH LINEAR	-0.00286446 -9.9443545E-05 0.00112667 -0.00023928	-0.10 -0.08 0.92 -0.72	0.9207 0.9359 0.3590 0.4755		0.02866021 0.00123186 0.00121976 0.00033344	•.		•

FULL HODEL FOR LOG OF PHOSPHATES

14:01 SATURDAY, NOVEMBER 14, 1981

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GEWERAL LINEAR MODELS PROCEDURE

LEAST SQUARES MEANS

PROB > 11 HO: LSHEAN(I)=LSHEAN(J) PROB > 111 SID ERR LNP

	0.0240 0.0168
HO: LSH	0.4146
PROB > 171	2 0.4146 3 0.0240
PROB > 111 H0:LSHEAN=0	0.2613 0.9704 0.0001
STD ERR LSMEAN	0.06545252 0.09928908 0.05451940
LSMEAN	0.07415940 0.00369804 0.24165080
ಕ	080

HOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISONS SHOULD BE USED. 0000 0000 5 -100

201	LSHEAN	LSMEAN	HO:LSNEAN=0		1/1 1 2 2 3 3 1 401 1/1 5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	3	SHLAN(U)	ŝ	
.	0.01278890 0.14379792 0.19686402 0.03102889 0.03102889	0.10159980 0.20890103 0.114281001 0.10067651 0.07563328	0.9002 0.4936 0.1727 0.1789 0.1604	2 0.5619 3 0.2905 4 0.8923 5 0.4002	0.5619	0.2905 0.6343 0.3373 0.5087	0.8923 0.6368 0.3373 0.3373	0.4002 0.6681 0.5087 0.5087	0.23
70	0.14715864	0.04502133	0.0017	6 0.168		-	0.2946	0.6439	•

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NOTE: TO EMSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISOMS SHOULD BE USED.

PROB > 171 HO: LSHEAN(1)=LSHEAN(J) 1/J 1 2 3	1 . 0.0815 0.0583 2 0.0815 . 0.9327 3 0.0583 0.9327 .
PROB > 111 HO:LSHEAN=0	0.9326 0.0054 0.0112
STD ERR LSHEAN	0.09734747 0.05608899 0.06370308
LSHEAN	-0.00826803 0.16144727 0.16632900
PINBL	- 0 10

NOTE: TO ENSURE OVERALL PROTECTION LEVEL, ONLY PROBABILITIES ASSOCIATED WITH PRE-PLANNED COMPARISOMS SHOULD BE USED.

HO:

PROB > 171 I LSMEAN1=LSNE	0.7696	PROB > IT 1 LSMEAN1=LSHE/	0.8248
PROB > 11! HO:LSMEAN=0	0.1082	PROB > 171 HO:LSMEAN=0	0.0999
STD ERR LSMEAN	0.06054154	STD ERR LSMEAN	0.06752568
LSHEAN	0.09857828	LNP	0.11267045
DIST	m 19	STMK	o-

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REDUCED MODEL FOR LOG OF PHOSPWATES--WH AND STMK ONLY 14:01 SATURDAY, MOVEMBER 14, 1681 Model Includes test for Interaction Between HH and Stmk 14:01 Saturday, Movember 14, 1981

GENERAL LINEAR MODELS PROCEDURE

HODELS PROCEDURE

CLASS LEVEL INFORMATION

CLASS LEVELS VALUES

STNK 2 01

NUMBER OF OBSERVATIONS IN DATA SET = 95

2 NOTE: ALL DEPENDENT VARIABLES ARE CONSISTENT WITH RESPECT TO THE PRESENCE OR ABSENCE OF MISSING VALUES. HOMEVER, OMLY Observations in data set can be used in this Amalysis.

				LEAST SQUARES HEANS	LEAST S			
			0.00025806	•	0.3118	-1,02	-0.00026263	HH LINEAR
			STD ERROR OF ESTIMATE	STD	PR > T	T FOR HO: PARAMETER=D	ESTIMATE	PARAMETER
0.311	2.50	0.09822616 0.04068153 0.03863828		0.3974	1.11 0.72 0.99	0.04373832 0.02643148 0.03883828		STNK Hif HH [®] STNK
× 84	F VALUE	TYPE IV SS	5	PR > F	F VALUE	TYPE I SS	01	SOURCE
0.2116579		0.19818247				3.29238767	10 .	CORRECTED TOTAL
LNP MEA		STD DEV		0.03927629	0.039	3.18137960	5	ERROR
613.633	0.033717	0.4260	0.94	0.03700269	0.037	0.11100808		HODEL
2	R-SQUARE	PR > F	F VALUE	QUARE	MEAN SQUARE	SUM OF SQUARES	DF	SOURCE
							LNP	DEPENDENT VARIABLE: LNP
								•

PROB > 171 HO: LSHEAN1=LSMEAN2 0.1883

PROB > 11 H0:LSHEAN=0 0.0001 0.0001

STD ERR LSHEAN 0.02806086 0.03613032

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LSHEAN LSHEAN 0.22983203 0.16913695

1 2 2

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116

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